

# RELATIONSHIP OF MELILITE-BEARING ROCKS TO KIMBERLITE: THE SYSTEM AKERMANITE - CO<sub>2</sub>

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The notion that kimberlite is genetically related to melilite "basalt", attributed to H. C. Lewis (1897), was revived after olivine melilitite was discovered in close association with a kimberlite breccia (Ukhanov, 1963). Reports of melilite in kimberlite have been rare and then the identification of the mineral was dubious or based on the interpretation of alleged pseudomorphs. Both CO<sub>2</sub> and H<sub>2</sub>O play an important role in kimberlite groundmass (olivine + phlogopite + calcite); however, akermanite (Ak), the principal component of melilite, was found to be stable only at water pressures below 10.2 kbar (Yoder, 1968). Because the stability field of phlogopite enlarges with CO<sub>2</sub>, akermanite was investigated in the presence of CO<sub>2</sub>.

Two mixtures having the composition 2CaO:MgO:2SiO<sub>2</sub>:CO<sub>2</sub> (= Ak + CO<sub>2</sub>) and 3CaO:MgO:2SiO<sub>2</sub>:2CO<sub>2</sub> (= Merwinite + 2CO<sub>2</sub>) were studied up to 10 kbar total pressure at temperatures from 950° to 1475°C. The results of the first composition run in gas-media, high-pressure apparatus for times up to 1 week are presented in Fig. 1.

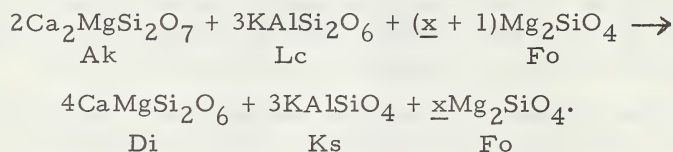
The maximum stability limit of akermanite in the presence of excess CO<sub>2</sub> was found to be 6.1 kbar. At relatively lower temperatures and below that pressure, akermanite reacts with CO<sub>2</sub> to form diopside (Di) + calcite (Ct). The reaction takes place through a range of temperatures because of possible solid solutions in dolomite in calcite, akermanite in diopside, and a variety of solid solutions in akermanite. The reaction curve, extended through the measurements of Walter (1963), terminates at its junction with the breakdown curve of akermanite → wollastonite + monticellite (Yoder, 1968) near 685°C at a pressure less than 100 bars. The appearance of diopside in the Ak + CO<sub>2</sub> composition is due primarily to the solution of additional components in the gas phase (G).

The melting of akermanite in the presence of CO<sub>2</sub> appears to be congruent and produces liquids which quench to a clear glass. The lowering of the melting curve relative to the anhydrous melting curve of Kushiro (1964) is believed to be due to the solubility of CO<sub>2</sub> in the liquid. The melting of Di + Ct also yields a clear, homogeneous glass without quench products. It is concluded that carbonatite is not necessarily a product of immiscibility from all ultrabasic compositions. New preliminary data on the melting of calcite cleavage blocks in the absence of excess CO<sub>2</sub> and the melting of diopside in the presence of CO<sub>2</sub> (Eggler, 1973) are given in Fig. 2.

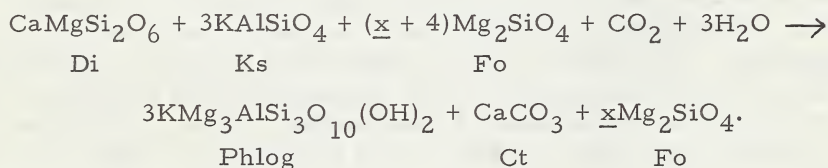
The restriction of akermanite to relatively low pressures in the presence of excess CO<sub>2</sub> is evident in Fig. 1. In the absence of free CO<sub>2</sub>, the assemblage Ak + Di + Ct, related to the alnoites, is stable

and would persist until the breakdown of Ak itself at about 15 kbar. Inclusion of the soda melilite component may extend this range to higher pressures; however, the minerals in natural melilite-bearing rocks were shown to react at high pressures to form pyroxenites in the absence of water and  $\text{CO}_2$  (Tilley and Yoder, 1968). Similar behavior would be expected in the absence of the additional phase calcite. It is not uncommon to see veins of calcite crossing pyroxenites without the formation of melilite. If kimberlite is derived from the mantle, then melilite-bearing rocks are not likely sources.

Melilite is commonly associated with leucite (Lc), kalsilite (Ks), and phlogopite (Phlog). Olivine melilite leucite (ugandite) is probably related to olivine kalsilite pyroxenite (mafurite) by the reaction:



Diopside + kalsilite occurs with melilite in metamorphosed sedimentary blocks at Brome Mountain, Quebec (Philpotts, Pattison, and Fox, 1967), and at Hendricksplaats, Bushveld (Willemse and Bensch, 1964). It is important to note that the three phases akermanite, monticellite, and leucite, which commonly occur together, all break down at relatively low pressures. Addition of  $\text{CO}_2$  and  $\text{H}_2\text{O}$  to different proportions of the above products yields the kimberlite groundmass assemblage:



It appears that there is some support for the suggestion that a magma having melilite affinities could be transformed into kimberlite with the aid of suitable volatiles. Loss of volatiles in transit or crystallization at low pressure would yield a melilite-bearing assemblage within the upper crust.

From the relations exhibited in Fig. 1 and the equations listed immediately above, it would seem that clinopyroxene ought to be a common residual phase in kimberlite. It is in fact a rare phase and then may be derived in part from fragmented xenoliths incorporated in the kimberlite. The large array of intermediate rocks that do contain clinopyroxene, melilite, and the major phases in the groundmass of kimberlite are usually described as alnoite.

The results presented in Fig. 1 are also of importance to the metamorphism of carbonaceous rocks. The reaction  $\text{Di} + \text{Ct} \rightarrow$

$\text{Ak} + \text{CO}_2$  is Bowen's (1940) step 8 in the metamorphism of a siliceous dolomite. Because of the exceptionally high temperatures involved where the  $\text{CO}_2$  pressure is the total pressure, it would appear that the reaction runs in nature at very low partial pressures of  $\text{CO}_2$ .

#### References

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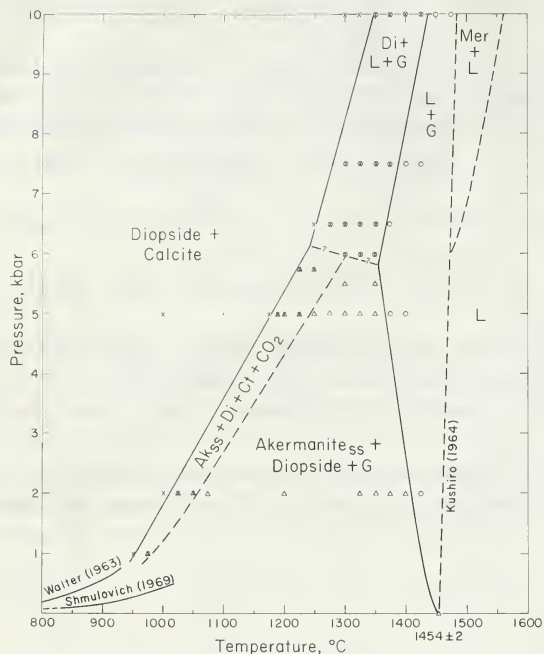


Fig. 1. Pressure-temperature diagram illustrating new results for akermanite- $\text{CO}_2$  using the diopside + calcite (1:1 mole) composition. Short-dash curves represent an extrapolation to low pressures of the melting data of Kushiho (1964) for akermanite.

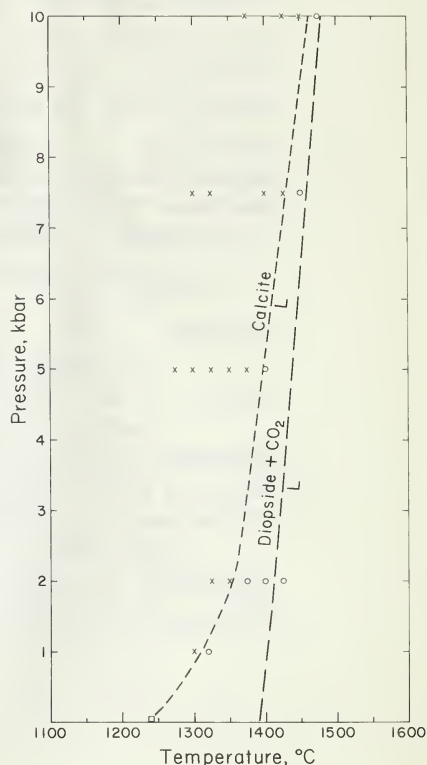


Fig. 2. Preliminary pressure-temperature curve for the melting of calcite cleavage blocks. Invariant point (square) from Baker (1962) and 1-kbar bracket from Wyllie and Tuttle (1960). The melting of calcite is compared with that of diopside in the presence of an excess of  $\text{CO}_2$  (Eggler, 1973). Calcite may melt at slightly lower temperatures in the presence of excess  $\text{CO}_2$ .