INCLUSIONS IN DIAMONDS: GARNET LHERZOLITE AND ECLOGITE ASSEMBLAGES. Martin Prinz

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Thirty inclusion-bearing diamonds have been studied in order to determine the mineral assemblages and chemical compositions of the inclusions. Inclusions were exposed either by cracking or burning the diamond or studied in <u>situ</u> within cut gems. They were analyzed with an electron microprobe, and supplemental data were obtained by optical goniometry and x-ray techniques. Inclusions range in size from <10 μ m to about 1 mm, but are mostly 50-200 μ m. Inclusions which are syngenetic with the formation of the diamond have good crystal faces, but they generally appear to have the morphology of negative crystals of diamond, rather than of their own crystal structure. Thus, garnets, pyroxenes, olivines, and other minerals are dominated by octahedral faces. Sobolev et al.(1972) has reported a similar finding. These crystals which have foreign faces, are termed xenohedral.

Inclusion assemblages fall into two suites, garnet lherzolite (or peridotite) and eclogite. Sobolev(1972) and Meyer and Boyd (1972) have found a similar grouping. Each has compositionally distinct mineralogy and therefore an assemblage can be assigned to one of the suites even when one mineral is present. Thus, the compositional range of assemblages reported here represents portions of assemblages from different diamonds.

Twelve diamonds contained portions of the garnet lherzolite assemblage which consists of garnet, diopside, enstatite, olivine, and chromite. Garnet is burgundy-colored, is pyrope-knorringite-rich, with 0.7-9.2% Cr₂O₃ (Table 1). Diopside has 0.11-1.56% Cr₂O₃ and 2.8-4.4% Al₂O₃. Diopside, in two diamonds, appears to have exsolved enstatite into wholly separate portions of a single crystal. Enstatite (En₈₆₋₉₄) has 0.06-0.28% Cr₂O₃ and 0.09-1.32% Al₂O₃ (Table 2). Olivine (F092-93), commonly present as elongated crystals, has 0.03-0.15 Cr₂O₃; the Cr²O₃ and 0.09 crystals. An about 66% Cr₂O₃; 6% Al₂O₃, and 10% MgO (Table 3).

Eighteen diamonds contained portions of the eclogite assemblage. The essential minerals are garnet and omphacitic pyroxene. Garnets are honey-brown in color and rich in pyrope, grossular, and almandine molecules (Table 1); they contain about 0.2% Na₂0. Omphacitic pyroxenes have 2.7-10.7% Al₂O₃, 1.4-5.8% Na₂O, and 0.04-0.87% K₂O (Table 2). K₂O contents of 0.62-0.87% were found in numerous crystals within 4 diamonds and are considerably higher than any previously recorded in omphacite; the K₂O is homogeneously distributed and x-ray data shows that no other phase² is intergrown. Other primary minerals present in the eclogite assemblage are kyanite(Table 3; reported for the first time in a diamond), olivine (in a rare olivine eclogite assemblage), rutile (Table 3),

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phlogopite, and magnetite. The phlogopite is in a polymineralic inclusion, with omphacite and rutile, is low in MgO and high in FeO and TiO₂(Table 3), and is the first recorded occurrence within a diamond. Magnetite is found in 6 eclogite assemblage diamonds and is a very pure phase (Table 3), except for some inhomogeneously distributed Ni and Co. Crystals are highly porous, and the origin of the porosity is not yet known. Its common occurrence in eclogite assemblage minerals suggests that it is syngenetic, although the oxygen fugacities necessary for its stability remain unknown. Sanidine, a polymorph of silica (possibly coesite), and a membranous material rich in SiO₂ Al₂O₂and K₂O are also present in diamonds with eclogite assemblage, but it² is not yet clear if they are syngenetic or epigenetic with the origin of the diamond.

Epigenetic minerals found in diamonds are muscovite, amphibole (hornblende, actinolite), sulfide (pentlandite, chalcopyite), and an altered $Si0_2$ -Al $_20_3$ -FeO phase which may have originally been kyanite.

The following results and conclusions are drawn from this study: (1) Diamonds appear to be a unique window into the upper mantle and provide unaltered samples of a wide compositional range of several rock suites. (2) No mixing of suites has been found within a single diamond. (3) Temperatures of formation of the pyroxenes in garnet Iherzolite assemblages are between 1000°C and 1400°C, depending upon whether the present compositions are used or the presumed compositions of the original homogeneous pyroxenes, assuming exsolution. The pressures of formation for the garnet lherzolite assemblages are at least 40 kb, but cannot yet be estimated with existing experimental data. (4) The presence of high (0.6-0.8%) K₀O in omphacitic pyroxene may be highly significant in terms of partial⁶melting models and heat sources in the upper mantle and should be studied experimentally. (5) Some kyanite eclogites are now known to be stable at temperatures and pressures at which diamond is stable. (6) The abundance and variety of inclusions, both monomineralic and polymineralic, within single diamond crystals suggests that a complex history has preceded the formation of the host diamond. This history would appear to precede the formation and emplacement of kimberlite and further study of inclusions in diamonds will reveal some of its details.

References:

Sobolev, N.V., Botkunov, A.I., Bakumenko, I.T., and Sobolev, V.S.(1972) Dokl. Akad. Nauk SSSR. <u>204</u>, 117-120.
Sobolev, N.V. (1972). Aust. Nat. Univ. Publ. No. <u>210</u>, 38 p.
Meyer, H.O.A. and Boyd, F.R. (1972). Geochim. Cosmochim. Acta <u>36</u>, 1255-1273. Table 1. Electron microprobe analyses of garnet inclusions in some diamonds (in weight percent).



Fig. 1. Ca-Mg-Fe plot of garnet and pyroxene data, and Forsterite-Fayalite (Fo-Fa) plot of olivine data from garnet lherzolite and eclogite inclusion assemblages in diamonds. Dashed lines connect coexisting minerals within a single diamond. Numbers refer to individual diamonds.

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Table 2. Electron microprobe analyses of pyroxene inclusions in some diamonds (in weight percent).

			Eclogite			
	Garnet Lherzolite Assemblage				Assemblage	
Diamond No.	3(Cpx)	3(0px)	l(Cpx)	1(0px)	38	42
Si0 ₂	55.6	58.1	55.5	57.3	55.8	54.7
Ti02	.23	.07	.14	.04	.19	.25
A1203	2.78	.09	4.4	1.32	10.7	7.6
Cr203	1.56	.11	.28	.06	.13	.13
V 20 3	.05	.01	.03	*	.09	.05
FeO	1.61	3.9	3.8	8.5	4.1	6.4
NiO	.03	.03	.02	.01	*	*
Mn O	.07	.06	.18	.21	.05	.11
MgO	15.6	35.7	15.6	31.4	9.8	10.4
CaO	19.5	.29	17.9	.63	13.3	15.6
Na ₂ 0	2.23	.08	1.82	.13	5.8	3.8
к ₂ õ	.11	*	.16	.02	.15	.87
Total	99.37	98.44	99.83	99.62	100.11	99.91
Kushiro (1962)	End Mem	bers				
NaAlSi206	16.3	. 4	13.1	. 9	40.9	30.9
CaAl ₂ SiO ₆	. 2	. 4	3.1	2.3	1.8	. 4
CaTiÃl ₂ 06	. 6		. 4	.1	.5	. 7
MgSiO3	42.5	88.7	43.2	83.7	26.1	28.2
FeSi03	2.6	9.9	6.2	13.0	6.3	9.9
CaSi03	37.8	. 6	34.0		24.4	29.9

Table 3. Electron microprobe analyses of selected minerals in some diamonds (in weight percent).

Garnet Lherzolite Assemblage			Eclogite Assemblage			
Diamond N	10.4b	51	30	18	18	.3.3
Mineral	Chromite	Kyanite	Rutile	Rutile	Phlogo- pite	Magne- tite
Si0 ₂	.11	36.6	.08	.24	36.6	.04
TiO ₂	.32	.13	90.0	96.7	10.8	*
A1 20 3	5.9	62.1	1.91	1.35	11.8	*
Cr_2O_3	66.2	.02	.17	.34	. 35	.09
V 20 3	.33	n.d.	.20	n.d.	n.d.	.01
Fe ₂ 0 ₃			7.2	1.14		68.4
FeO	16.1	.30			12.1	30.8
NiO	.14	n.d.	.21	n.d.	n.d.	. 38
MnO	.22	*	.11	.05	.04	.01
MgO	10.0	.04	.07	.01	12.6	*
CaO	.03	.01	. 42	.26	.09	.04
Na ₂ 0	n.d.	.03 -	n.d.	n.d.	.10	n.d.
к ₂ о	n.d.	*	n.d.	n.d.	9.6	n.d.
Total	99.35	99.24	100.37	100.09	94.08	99.92
** CoO	, 0.23%.	Ni and	Co varia	ble with	nin cryst	als.