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Tectonic setting of kimberlites

Hielke Jelsma¹, Wayne Barnett², Simon Richards³, and Gordon Lister³

¹De Beers Exploration, P/Bag X01, Southdale 2135, South Africa ²De Beers Mineral Resources Management, P/Bag X01, Southdale 2135, South Africa ³Research School of Earth Sciences, The Australian National University, Canberra, Australia

Kimberlites and related alkaline rocks (kimberlites *sl*: e.g., kimberlites, lamproites, melnoites, carbonatites) range in age from the Neoarchean to the Cenozoic, and are found on all continents and intrusive in a variety of geological domains (Fig. 1). Their spatial and temporal distribution has been related to mantle plumes or hotspots (Morgan, 1983; Le Roex, 1986), or to tectonic processes involving lithospheric faults formed or reactivated during rifting of continents (Dawson, 1970; Marsh, 1973; Sykes, 1978; Bailey 1992; White et al., 1995; Jelsma et al., 2004; Moore et al., 2008).



Fig.1 Cratonic domains (grey) and global distribution of kimberlites (red), lamproites (orange), melnoites (green), and carbonatites (blue)

In this study the tectonic setting is reviewed. Spatial and temporal data for kimberlites are compared with the lithosphere anisotropy and surface geological features. It is argued that kimberlites are records of tectonic events in a changing world, marking stages in the assembly or break-up of continents, and are key pieces in palaeocontinent reconstructions.

Temporal and spatial constraints

Over time, each continent shows alternating periods set apart either by abundance or scarcity of kimberlite magmatism. Table 1 and Fig. 2 provide an overview of available age data using Southern Africa as an example. Abundance peaks occur during the Mesoproterozoic (1650 Ma, 1330 Ma, 1195 Ma, 1100 Ma), the early Palaeozoic (520 Ma, Pan-African), the Mesozoic (240 Ma, 145 Ma, 121 Ma, 85 Ma, 73 Ma), and the Cenozoic (54 Ma and 31 Ma). Some kimberlite groups are better preserved compared to others, depending on the geological environment (basically an exhumation and burial story: 1km erosion off-the-top of a kimberlite pipe takes out most of the crater- and diatreme components).





Fig. 2 Kimberlite age data for Southern Africa

In space, Southern Africa is transected by a network of systematic continental lineaments that are considered to be corridors of concentrated, aligned tectonic activity (Fig. 3). Some of these are lithosphere-scale structures that formed focal areas for (often repeated) kimberlite magmatism representing anisotropies associated with gradients in the subcontinental lithosphere mantle (SCLM), measured by seismic velocity, resistivity, or compositional heterogeneities. Within the crust, these gradients may be manifest by (a) terrane boundaries or marginal orogenic belts, (b) incipient continental rifts, (c) extensions of oceanic fracture zones (FZ) or associated accommodation zones within continents, or (d) major dyke swarms. On a regional scale, the distribution of kimberlites is typically clustered along these trans-lithospheric (through-going structural corridors fundamental "flaws" within the continental architecture); many have been repeatedly reactivated over time.



Fig. 3 Distribution of kimberlites in Southern Africa, kimberlite age data, and structural elements, mapped from geological, geophysical, remotely sensed and elevation data

In Southern Africa, a strong association is noted between Mesozoic kimberlites and NE or NW trending structures (e.g., Jelsma et al., 2004). Similar lithosphere trends are observed as anisotropies associated with the cratonic keel (James et al., 2001), craton edges and marginal orogenic belts, the Karoo event, the East Africa Rift, as well as the extension of oceanic FZs into the African continent (Fig. 3, cf. Dirks et al., 2003). Both are associated with Gondwana break-up (see marked fault offsets along the continent margins of Africa, e.g. in Somalia-Kenya, Angola-Namibia).

On a local scale, a key feature is a preferred common long axis of kimberlite pipes or orientation of kimberlite dykes amongst coeval kimberlites (Fig.4). This is observed within a cluster or across several clusters within a continental domain and implies a regional tectonic control on emplacement.



Fig. 4 Lesotho kimberlite dykes (width map 100 km), 90 Ma

Bottom-up or top-down?

The Jurassic-Cretaceous kimberlites in Southern Africa are clustered along sets of parallel lineaments widely distributed across a large part of the subcontinent, with the younger groups found closer towards the continent margins (Fig. 3). Repeated magmatism is seen within a number of clusters. Parallelism of same-age kimberlite "tracks" and repeated magmatism at the same sites suggests intrusion along structural discontinuities as a result of plate tectonic processes (Anderson 2001; Vaughan and Scarrow, 2003) rather than an association with deep-seated plumes and hotspot tracks. Systematic younging of kimberlites between 70 Ma (E) and 46 Ma (W) in the western part of South Africa and in Namibia may be attributed to stress propagation.

Kimberlites	Events	Plate motion vector	Offshore sediments	Onshore sediments	Hotspot manifestation
"sensu lato"		ausps	unconformities	unconformiões	thermal insulation
??	183 Ma: Karoo, onset spreading				183 Ma: Bouvet (Karoo-Ferrar)
148-141 Ma	Weddell/Somall Basins 145 Ma: W-E Gondwana separate	150-142 Ma: AFR-EURAS		142 Ma: major folding event	
135-130 Ma 127-114 Ma	133 Ma: Etendeke-Parana 133-120 Ma: progressive opening of South Atlantic	135-133 Ma: AFR-ANT 130-125 Ma: AFR-ANT	127-124 Ma: 541		132 Ma: Tristan, Shona, Kerguelen
	120 Ma: AFR-SAM and IND-ANT separate 118 Ma: Lower Cretaceous Rift System (C-AFR)	121 Ma: AFR-EURAS	120-118 Ma: 6 Af	120 Ma: tolding event	
??			112 Ma: 13A/J		
103-101 Ma				101 Ma: folding event	
05 00 14					
95-83 Ma	85 Ma: IND-MAD separate	84 Ma: AFR-EURAS	94-91 Ma: 1540 86-84 Ma: 1640	84 Ma: major folding event	85 Ma: Marion, St Helena, 90East, Compress, Reunion
76-70 Ma	72 Ma: IND-SEY separate		74-67 Ma: 21-2241		Contraria, reserven
65-59 Ma		65 Ma: AFR-EURAS		66 Ma: major folding event	
56-50 Ma ?? 32-31 Ma		37 Ma: AFR-EURAS	54-43 Ma 40-37 Ma	37 Ma: tolding event	
Table 2 - age data	Jelsma et al., 2004	Fairhead and Wilson 2004, Eagles and König 2008	McMillan and Dale 2002	Fairhead and Wilson 2004	www.mantleplumes.com

Fig. 5 Time-event chart for Southern Africa

Fig. 5 compares the timing of magmatism with major events affecting Gondwana. Kimberlite emplacement



age windows coincide with: (i) cusps and jogs in the relative plate motion paths of continents (AFR-EURAS, AFR-ANT), based on a fit of seafloor magnetic anomaly isochrons with oceanic FZ traces or on the construction of apparent polar wander paths from palaeomagnetic data; (ii) unconformities in the offshore and onshore stratigraphic record; (c) hotspot family manifestation. Hargraves and Onstott (1980) already noted the synchronicity of kimberlite ages and changes in Apparent Polar Wander curves. In Hargraves words, "these relationships suggest that the emplacement of kimberlites may coincide with episodes of changes in the direction of plate motions."

Kimberlite age data for "Laurasia" and West Gondwana are shown in Fig. 6. Members within each of these supercontinents (e.g. Southern Africa and South America in West Gondwana) show similar age peaks, reflecting a common history, but different peaks when members across are visible the two supercontinents are compared. For instance, the 370 Ma reconstruction shows continent break-up and associated magmatism affecting Siberia, and continent stability and lack of magmatism in Southern Africa and South America. In North America Jurassic-Cretaceous kimberlite magmatism begins following the opening of the Central Atlantic and CAMP magmatism at 201 Ma. In West Gondwana, magmatism begins following the onset of spreading in the Weddell Sea and Somali Basin at 155 Ma. At 124 Ma, seafloor spreading in the South Atlantic causes rifting of Africa and South America and is accompanied by a major peak in kimberlite magmatism. The subsequent peak at 85 Ma is marked by concomitant magmatism in Southern Africa and South America, after continental break-up, and heralds major plate reorganization (Fig. 5).

The temporal association between the main pulses of kimberlite magmatism and hotspots (starting at 183 Ma, 130 Ma, and 85 Ma) and associated Large Igneous Province (LIP) magmatism is also significant. Sears (2001) proposed that the stalled Gondwana supercontinent "drove its own break-up by insulating the underlying mantle, causing thermal expansion, uplift, fracturing and associated LIP magmatism".

The trigger for kimberlite magmatism appears to be related to tectonic change - stress change induced within the continental plate because of supercontinent fragmentation (or late orogenic collapse, 520 Ma), and changes in the direction and velocity of plate motion (e.g. Marsh, 1973; Jelsma et al., 2004; Moore et al., 2008). Unconformities in the stratigraphic record in onshore and offshore basins document this episodic tectonic instability (Fig. 5; cf. Jelsma and Smith, 2004; Moore et al., 2008). Strain is accommodated on discontinuities within the lithosphere causing localized decompression melting near the base of the SCLM (Vaughan and Scarrow, 2003). In contrast, periods devoid of kimberlite magmatism correspond to times of supercontinent stability (a sluggish Gondwana at 500-250 Ma), or smooth plate motion paths.

Variable endowment across continental regions may be related to the transient chemical composition and physical state of a potential kimberlite source area, plus a required minimum size of continents to be able to cause effective insulation of the underlying mantle, plus the presence and orientation of anisotropies within continents to be exploited as transfer faults and to be able to allow effective stress propagation within the continents, plus tectonic change.



Fig 6 Kimberlite age data for "Laurasia" (top) and West Gondwana (bottom) with continent reconstructions at 370 Ma and 170 Ma (Torsvik, pers. comm., 2003)

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Table 1 Age data for kimberlites in Southern Africa

