

DIAMONDIFEROUS KIMBERLITES ON VICTORIA ISLAND, CANADA: A NORTHERN EXTENSION OF THE SLAVE CRATON

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INTRODUCTION

Victoria Island, located in the Northwest Territories and Nunavut in the High Arctic of Canada (Figure 1) has been a diamond exploration target since 1993 when exploration companies recognized the possibility Archean age rocks beneath the eastern half of Victoria Island under thin Paleozoic carbonate cover. Much evidence has been collected through work since 1993 that suggests Victoria Island has the potential to become a world-class diamond-producing district.

Recent exploration efforts have identified kimberlites with high diamond counts and robust diamond size distributions trending toward large 1 to 2 mm stones from small kimberlite samples. Supportive mineral chemistry including major oxides and trace element analyses suggests a cool geotherm, presence of potential peridotitic diamond sources and a strong diamond bearing eclogite signature.

GEOLOGY

Archean rocks exist in the eastern half of Victoria Island. Paleozoic sediments including dolomite, minor limestone, sandstone and shale of the Arctic Platform have been preserved within post-depositional structural basins that are bounded by the Precambrian Minto, Wellington and Duke of York inliers. Granitic rocks within the Wellington Inlier, east of the Victoria Island kimberlite field, yielded a U-Pb Zircon age of $2,601 \pm 3$ Ma, confirming that the basement is a northern extension of the Archean Slave Province (LeCheminant, 1996). Granodiorite dated at 2.4 Ga (Thorsteinsson and Tozer, 1962) in the southwest end of Hadley Bay intrudes metasediments that are by association Archean in age.

A broad syncline trending northeast across Victoria Island from Amundsen Gulf to Hadley Bay is composed of Precambrian strata of the Shaler Group. Approximately 2700m of sandstone, siltstone, shale, limestone, dolomite, and gypsum underlie 300m of

basalt (Natkusiak Formation 723 Ma) and agglomerate which make up the Shaler Mountains that are intruded by gabbro sills. The lower sandstone and gabbro units of the Shaler Group are also found along the southern coast of Victoria Island, and as small inliers south of Washburn Lake.

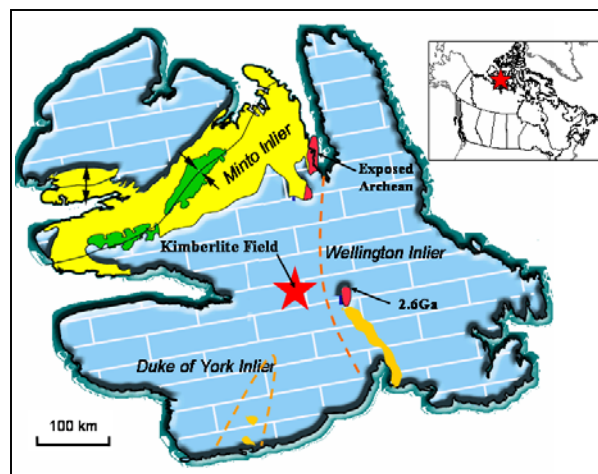


Figure 1: Location and geology map for Victoria Island.

Regional magnetics indicate the basement rocks have been intruded by NW-SE trending Mackenzie (1270 Ma) and Franklin (720 Ma) diabase dykes and a previously undocumented north-south trending diabase dyke system dated at 425 Ma (Diamonds North Internal Report, 1999). Upper Cambrian to Lower Devonian sediments of the Arctic Platform covers most of the island. Their presence limits the emplacement age of kimberlites exposed beneath glacial overburden to less than 385 Ma. Victoria Island kimberlites have been dated at 256-286 Ma (Diamonds North Internal Report, 2000).

KIMBERLITES

On Victoria Island kimberlites occur as distinct pipes, dykes and as blows along several dyke structures. A semi-continuous string of kimberlite, known as the Galaxy Trend, has a strike length of 20 kilometers. Geophysics has been an effective tool in the discovery of kimberlite on Victoria Island. The kimberlites discovered to date display distinctive,

discrete magnetic signatures, both magnetically positive and negative.

The non-magnetic nature of the 100m to 200m thick carbonate cover reduces the noisy magnetic signature of the Archean basement, allowing the younger kimberlites to stand out prominently with much higher frequency signatures. A typical example, the Snow Goose kimberlite, is illustrated in Figure 2.

The kimberlites of Victoria Island are typically rich in indicator minerals. Surficial sediment sampling is an effective means of exploration for these bodies. Indicator mineral grains include garnet, ilmenite, chromite and Cr-diopside.

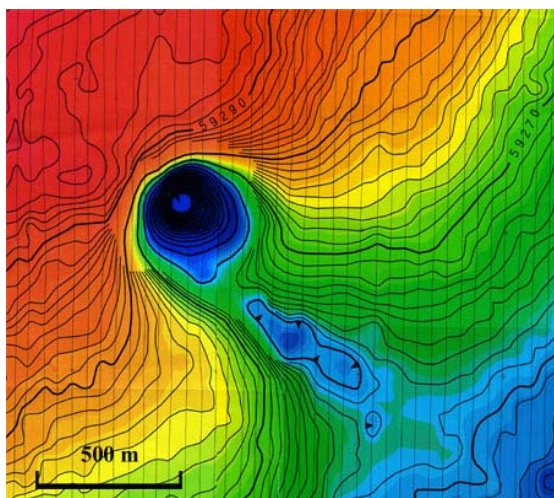


Figure 2: Magnetic signature over the Snow Goose kimberlite.

THERMOBAROMETRY

Recent trace element analyses for peridotitic garnets from the King Eider and Juno kimberlites indicate a cool geotherm of 38 mW/m² at the time of kimberlite emplacement for the central part of Victoria Island. The geotherm is constrained by the garnet P-T values with the lowest temperatures in the dataset, depicted in Figure 3. This geotherm is particularly important for Victoria Island as it indicates the thermal effects of the Mackenzie plume on the underlying lithospheric mantle were negligible. The 1270 Ma Mackenzie event is much older than the alkali-volcanism responsible for Victoria Island kimberlite emplacement. The recovery of diamonds from these kimberlites negates detrimental thermal effects of the Mackenzie plume.

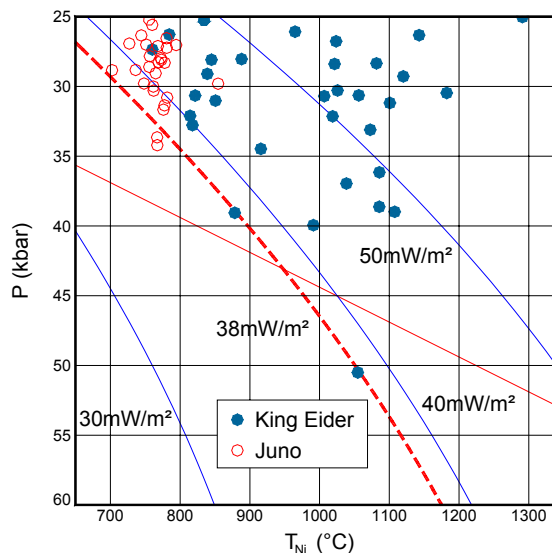


Figure 3: Geotherm of 38mW/m² estimated by G. Read using data from King Eider and Juno Kimberlites. After Griffin, 1989.

Along a 38mW/m² geotherm the diamond stability field is entered at 950°C and the base of the lithosphere, defined by the Yttrium edge, is at 1250°C. Therefore, the mantle beneath Victoria Island has a diamond window between 950°C (145km) and 1250°C (200km). For selected Victoria Island kimberlites, the garnet Ni-temperature histogram depicted in Figure 4 indicates the diamond window was well sampled as more than 80% of the garnets fall within the optimal temperature range.

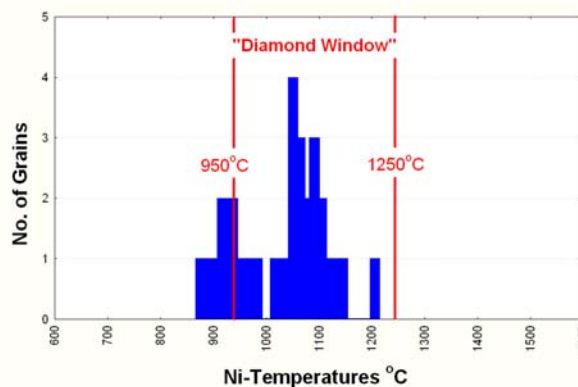


Figure 4: Histogram of garnet Ni-temperatures for King Eider Kimberlite. Ni-temperatures estimated using calculation by Canil 1999.

MANTLE COMPOSITION

Mineral abundances and compositions for several kimberlites on Victoria Island indicate the lithospheric mantle underlying Victoria Island is

comprised of both peridotitic and eclogitic parageneses. With respect to the peridotitic paragenesis, The CaO vs. Cr₂O₃ plot (Figure 5) demonstrates that the peridotitic paragenesis includes garnet harzburgite, and high Cr₂O₃ lherzolite (>10 wt% Cr₂O₃). These peridotitic populations suggest that the host kimberlites have the potential to carry diamonds. A group of very subcalcic garnets between approximately 2% and 6% Cr₂O₃ are also of interest.

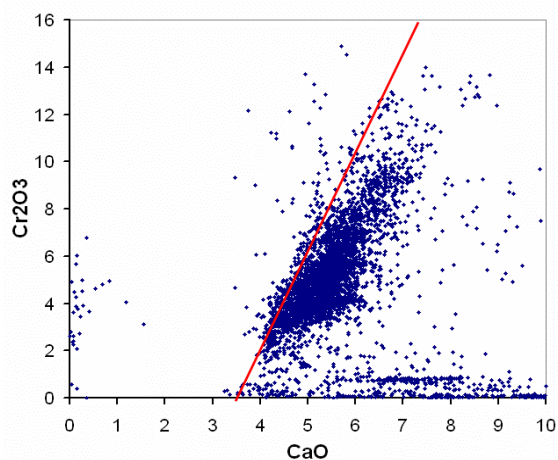


Figure 5: CaO vs. Cr₂O₃ plot of peridotitic garnets from several kimberlites on Victoria Island.

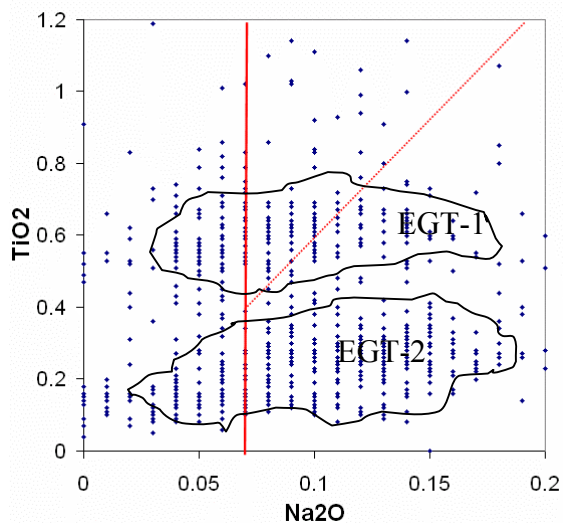


Figure 6: Diagram showing two populations of low Cr₂O₃ garnets.

The garnet population (Figure 5) contains low Cr₂O₃ garnets. These grains are further discriminated on a Na₂O vs. TiO₂ diagram (Figure 6), which indicates a high abundance of Group I eclogite garnets, and very

few megacrystic compositions. The eclogitic garnets have been divided into two sub-groups, EGT-1 and EGT-2. The Cr₂O₃ contents of the two distinct garnet populations are illustrated by the peridotitic plot (Figure 5) and the histogram in Figure 7. Elevated Cr₂O₃ and higher TiO₂ characterize the EGT-1 garnets. These garnets are possibly derived from websteritic source rocks. Similar websteritic garnet compositions are seen at the Victor and AT-56 kimberlites in the James Bay Lowlands of Ontario, Canada and are potential diamond source rocks for these kimberlites (Armstrong et al. 2003). EGT-2 garnets contain lower Cr₂O₃ and TiO₂ and represent typical Group I eclogite compositions that may be associated with diamonds.

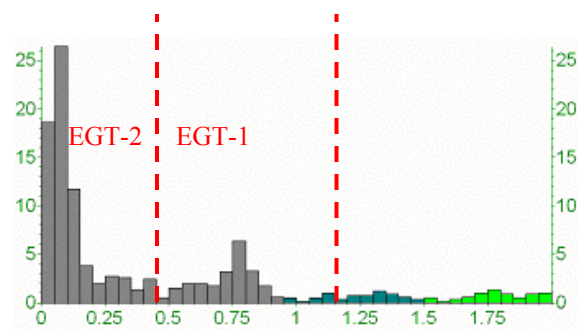


Figure 7: Histogram showing Cr₂O₃ content in Low Cr₂O₃ garnets – EGT-1 and EGT-2 groups.

DIAMOND PRESERVATION

Ilmenite grains recovered from Victoria Island kimberlites indicate that the rising kimberlite magma was under reducing conditions resulting in favourable diamond preservation. These data correspond with observations of diamond morphology. In general, many diamonds have distinct crystal faces, exhibiting little resorption, and are 85% to 95% preserved.

The ilmenite compositions from Victoria Island kimberlites are shown in different colours on Figure 8. All kimberlites are represented on the reduced side of the graph, and the distribution of and variations in ilmenite compositions may reflect mantle provenance.

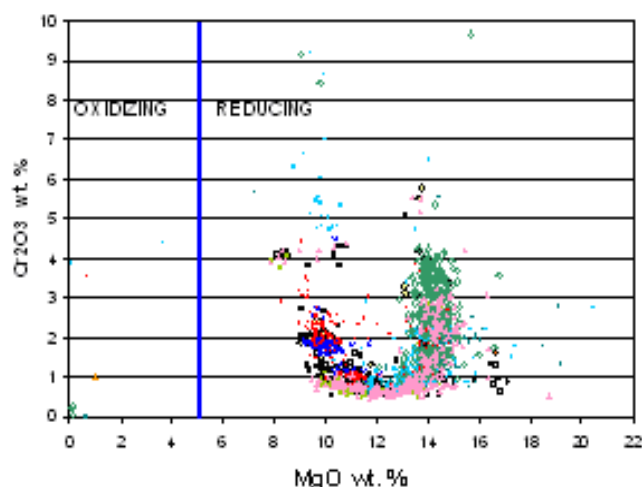


Figure 8: Kimberlitic ilmenite compositions for several kimberlites on Victoria Island.

LITHOSPHERIC MODEL

Kimberlite xenocryst paragenesis, kimberlite and basement geochronology and the presence of diamonds enable the development of a speculative lithospheric model for Victoria Island (figure 9). Archean age basement rocks underlying the Paleozoic sediments suggest likely northward extension of the Slave Province. This is supported by a cool cratonic geotherm calculated from garnet Ni-thermometry. The proportion of garnets with temperatures in the diamond window suggests substantial sampling by kimberlites at depth between 145 and 200 km. Kimberlite xenocryst composition suggests the sampling of peridotite (harzburgite, lherzolite and websterite) and eclogite at mantle depths within the diamond stability field. The abundance of eclogitic garnets present in the kimberlites suggests significant eclogitic component to the lithosphere beneath Victoria Island.

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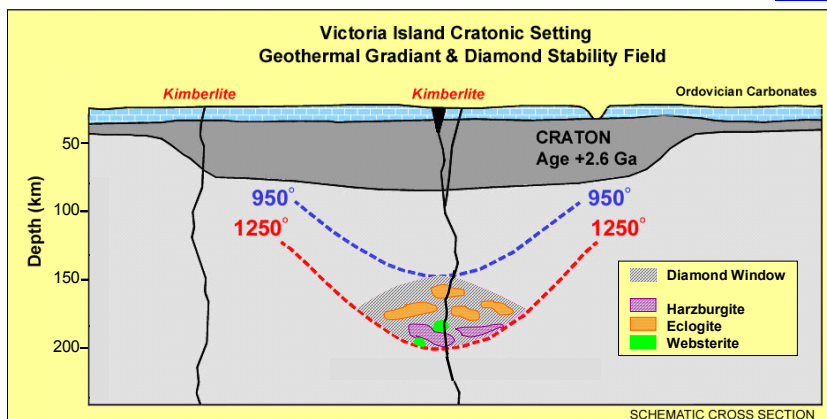


Figure 9: Speculative lithospheric model for Victoria Island.