ON THE SEARCH FOR NEODYMIUM-142 IN TERRESTRIAL ROCKS

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Abstract:

Neodymium-¹⁴²Nd is a decay product of ¹⁴⁶Sm (half-life of 103 Myr). ¹⁴⁶Sm is a so-called extinct radioisotope[1]. It existed 4.56 Ga ago $({}^{146}Sm/{}^{144}Sm)$: 0.008 ± 0.0011 [2]) but was essentially extinct after 300Myr. There is indirect evidence for the existence of ¹⁴⁶Sm by observation of ¹⁴²Nd variations in meteorites [1]. The first finding of ¹⁴²Nd in terrestrial rocks was reported from a single sample from the Isua supercrustals [3]. Several groups challenged this finding [4]. First attempts to find a ¹⁴²Nd anomaly on Earth was simply guided by measuring old rocks hoping that some of the Isua rocks are derived from an early differentiated reservoir. We found reasons to believe that some of the cratonic lithosphere might actually be a remnant of a primary differentiation. The first indication in this direction was the finding that the cratonic ultramafic in South Africa are having a higher modal abundance of orthopyroxene [5] than the primitive mantle. Later, the higher modal orthopyroxene contained was also demonstrated for the Siberian craton [6]. Herzberg actually proposed that this might be caused by primary differentiation on Earth [7]. We found that besides other geochemical peculiarities, a depletion of Al is

evident in such cratonic ultramafics. A geochemical parallelism is found in SNC meteorites and urelites. In SNC meteorites, variation in ¹⁴²Nd is evident [8] indicating that the Al deficit was caused by a very early differentiation. similar style primary А of differentiation could have produced the cratonic lithosphere. This made us search for a ¹⁴²Nd anomaly in cratonic ultramafics. In one of our samples we have found a $+30\pm10$ ppm deviation from the standard ^{142Nd}/¹⁴⁴Nd ratio. However, this needs further confirmation additional by measurements

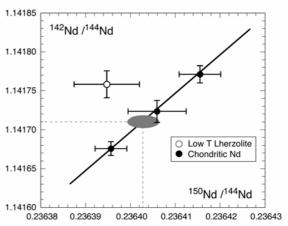
Introduction:

The formation of the solar system was very fast. The oldest solid fragments condensed at 4.571Ga, while at 4.557Ga (only 14 my later) basaltic generated magmas were on planetesimals, which in a short time formed the planets in the solar system. At 4.571Ga a number of radioactive isotopes existed which are extinct by now. In cosmochemisry we are using the decay products of such extinct isotopes in order to date the early evolution of the solar system: ²⁶Al – ^{26}Mg , $^{53}Mn - ^{53}Cr$, $^{182}Hf - ^{182}W$, $^{146}Sm - ^{142}Nd$, $^{129}I - ^{129}Xe$ are some of the frequently used isotopic systems. From studies of the 182 Hf – 182 W system there is the clear conclusion that at 4.541Ga the formation of the Earth was completed. At this time a number of short-lived radioactive isotopes still existed. ¹⁴⁶Sm was present at this time (¹⁴⁶Sm/¹⁴⁴Sm - 0,006). In a reservoir with a chondritic Sm/Nd ratio the 142 Nd will have today the normal ¹⁴²Nd/¹⁴⁴Nd ratio, of 1.141733; while at 4.550 Ga in the early Earth this ratio was 1.141233. If we find a reservoir which was depleted earlier in Sm we expect to measure a lower than chondritic ¹⁴²Nd/¹⁴⁴Nd. This effect is limited and as lower limit might be by a total depletion of Sm $(^{142}Nd/^{144}Nd)$ = 1.141233). On the other hand, in a reservoir where the Sm/Nd ratio is higher than chondritic (light rare earth depleted reservoir) the ¹⁴²Nd/¹⁴⁴Nd will higher than be chondritic. The enrichment of ¹⁴²Nd is only limited by the degree of early REE depletion (Sm/Nd ratio). However. this enrichment of ¹⁴²Nd will only occur so as long as 146Sm was still present (300myr after formation of the Earth.) However, the enriched ¹⁴²Nd signature of the early-depleted reservoir will remain practically forever no matter what differentiation processes overprint the reservoir. The only way to destroy this enriched ¹⁴²Nd is to remix the early-enriched reservoir with the earlydepleted reservoir. That means that lithologies involved the early in differentiation of the Earth should possess а "marked" Nd isotopic composition even today.

Experimental:

We developed a method to measure high-precision ¹⁴²Nd/¹⁴⁴Nd ratios in order to measure SNC meteorites which are believed to represent rock fragments from planet Mars. Those meteorites are very expensive and some of these meteorites are only 20g as a total sample. Therefore we had to develop a new method using only small samples. For this reason we measure Nd as NdO^+ ions in the mass spectrometer. While the measurement of NdO^+ is more sensitive, it has an additional correction (oxygen fractionation). which limits the precision of the results.

To eliminate all isobaric interferences (especially ¹⁴²Ce) we established a special chemical extraction routine to isolate Nd. However the challenge was to find a mass pectromeric routine to get the reproducibility to 7 ppm and the precision to 10ppm. The "second order mass fractionation" was careful analysed and we where able to eliminate the causes of this source of error.



Results:

We measured a number of SNC meteorites and got a very good agreement in those cases where ¹⁴²Nd where already measured by other groups. These results will be published elsewhere. However, we also made an attempt to measure ¹⁴²Nd on terrestrial samples. The measurement of ¹⁴²Nd is very time consuming, taking about 3 weeks of work for one sample.

So far we have found two ultramafic xenoliths (low temperature lherzolites) out of five which have a ¹⁴²Nd anomaly. The final deviation from the chondritic ¹⁴²Nd is shown in the diagram. The fractionation of oxygen in the masspec is a serious problem in our setup as mentioned above. In order to quantify the influence of oxygen fractionation we plot ¹⁴²Nd versus ¹⁵⁰Nd both normalized to ¹⁴⁴Nd and to ¹⁴⁶Nd/¹⁴⁴Nd = 0.7219. ¹⁵⁰Nd is only affected by oxygen fractionation and for chondritic ¹⁴²Nd we find a correlation between ¹⁴²Nd and ¹⁵⁰Nd. Most of the standard runs are plotting within the grev oval on this correlation line, while some of the extreme fractionated standard runs plotted along are the oxygen The fractionation line. two low temperature lherzolites, however, plot of this fractionation line because of elevated ¹⁴²Nd.

Discussion and conclusions:

This results indicate that the cratonic lithosphere contain remnants of the primary differentiation of the earth. We selected this type of sample since to our

opinion there were already signs of a early differentiation in verv the chemical composition of the cratonic lithosphere. The first indication in this direction was the finding that the cratonic ultramafics in South Africa have a higher modal abundance of orthopyroxene [5] than the primitive mantle. higher Later. modal orthopyroxene was also showenfor the Siberian craton [6]. Herzberg actually proposed that this might be caused by a primary differentiation on Earth [7]. We found that. besides other geochemical peculiarities, a depletion of Al is evident in such cratonic ultramafics. A geochemical parallelism can be found in SNC meteorites.

References: [1] Lugmair and Marti (1977) EPSL,

35, 273–284. [2] Lugmair and Galer (1992) GCA 56,1673–1694. [3] Harper and Jacobsen (1992) Nature 360, 728– 732. [4]Goldstein S L and Galer SJG (1992)Eos P323; Sharma et al. (1996) GCA 60,2037–2047. [5] Boyd F. R. (1989) EPSL 96, 15–26.[6] Boyd F. R. et al. (1997) Cont. Min. Pet., 128,228– 246. [7] Herzberg C.T. (1993) EPSL 120,13–29. [8] Harper et al. (1995) Science 267, 213–217.