Trace Element and Isotope Geochemistry of Siberian Kimberlites

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Introduction

As part of a larger study investigating the Hf-Nd isotope compositions of kimberlites and lamproites from South Africa, Australia, and North America (Nowell et al. 1998a, b), we have undertaken a geochemical survey of a small number of kimberlites from Siberia. The areas under investigation include the Malo-Botuobinsk field (Mir, Internationalaya) and the Daldyn-Alakit field (Udachnaya, Aikhal, Jubilee, Sytikanska and Zarnitsa); the Udachnaya sample is from the East pipe. Samples were collected from open pit workings where possible, or from dumps, and are mostly tuffisitic kimberlite breccias (TKB's) containing relatively abundant megacrysts of magnesian-ilmenite, chrome-pyrope and altered olivine. Inclusions of mega-crystalline dunites are also common in some samples. The freshest possible material was obtained, with areas of obvious nontronite alteration being avoided. Significant effort was expended in picking out as much of the xenolith and megacrystic material as possible during sample crushing.

Geochemical Results

Results of trace element and isotopic analysis are presented in Table 1 and Figures 1-3. In extended trace element diagrams (i.e. 'spider' plots) the Siberian kimberlites lie within the range of Group I kimberlites from South Africa, showing strong enrichment of the LREE as well as relative depletions in Sr, Hf and Zr (Fig. 1). Two samples from the Aikhal pipe (A3 and A4) have virtually identical trace element patterns, except that A4 has significantly lower Pb contents than A3. Sample M3 has a slightly different pattern for the most incompatible LIL elements when compared with the other Siberian samples in that it lacks a strong Ba spike.

87Sr/86Sr In a plot VS. 143Nd/144Nd (Fig. Siberian samples define a small cluster of points that lie just outside the field of Group I kimberlites from South Africa, having slightly higher 87Sr/86Sr at a given ¹⁴³Nd/¹⁴⁴Nd. are clearly distinct from Transitional and Group kimberlites.

In contrast to their Sr-Nd isotope systematics, Pb isotopes for the Siberian kimberlites show a very wide range. Three samples lie within the field of

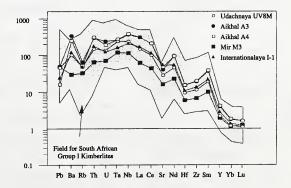


Figure 1. Spider plot of Siberian kimberlites compared with Group I kimberlites from South Africa. Concentrations are normalized to primitive mantle.

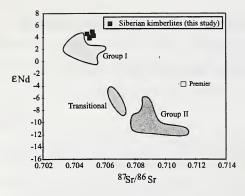
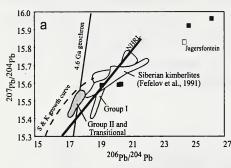


Figure 2. Plot of Sr vs. Nd isotopes comparing Siberian and South African kimberlites.

Group I kimberlites, but 2 extend to some of the highest ²⁰⁶Pb/²⁰⁴Pb, ²⁰⁷Pb/²⁰⁴Pb and ²⁰⁸Pb/²⁰⁴Pb values known for kimberlites. This is consistent with data for similar Siberian kimberlites (Fig. 3) described by Fefelov et al. (1991), who also observed that the higher ratios occur in the most carbonatized samples. Note that even though the two Aikhal samples analyzed in this study have nearly identical trace element compositions, they have vastly different Pb isotope compositions (Table 1), almost certainly acquired during or postemplacement.

Most of the Siberian kimberlites intruded evaporites and/or brines that are known to contain veins of calcite, sulfide and bitumen. Bitumen is known to have very radiogenic Pb isotope compositions because it typically contains disseminated uraninite (Parnell and Swainbank, 1990). Given the high concentrations of U (>1 wt%) and extremely radiogenic Pb compositions of some bitumen (206 Pb/ 204 Pb >100), it is possible that the high ratios observed in the Siberian kimberlites are a function of contamination at relatively shallow levels in the crust during emplacement. This contamination appears not to have affected the Sr-Nd isotope or trace element compositions of the rocks, presumably because the contaminant was relatively low in these other trace elements.



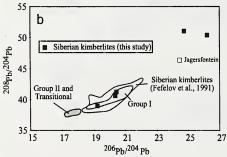


Figure 3. Plot of $^{206}\text{Pb/}^{204}\text{Pb}$ vs $^{207}\text{Pb/}^{204}\text{Pb}$ (a) and $^{208}\text{Pb/}^{204}\text{Pb}$ (b) comparing Siberian and South African kimberlites.

Conclusions

Siberian kimberlites are similar to Group I kimberlites from South Africa. Pb isotope systematics have probably been disturbed by shallow level contamination processes related to emplacement in the evaporite country rock.

References

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Table 1: Trace element (ppm) and Sr-Nd-Pb isotope data for Siberian kimberlites

Sample	UV8M	A3	A4	M3	I-1
Sc	12.33	16.19	16.82	10.15	11.77
Ti	1.37	0.51	0.52	1.31	0.48
\mathbf{V}	105.70	71.69	66.42	92.93	60.16
Cr	1437.52	1715.46	1139.73	870.24	1571.60
Mn	0.12	0.10	0.10	0.11	0.12
Co	76.75	65.18	53.61	94.54	56.37
Ni	1140.30	1023.07	1055.12	1557.72	1486.39
Cu	59.01	25.49	18.97	38.22	25.06
Zn	33.44	43.21	8.79	44.10	29.53
Ga	4.93	5.58	5.35	4.10	4.63
Rb	30.43	41.04	33.03	20.69	30.04
Sr	593.01	848.77	993.73	339.24	1131.29
Y	10.09	18.17	19.14	8.72	12.91
Zr	126.58	216.16	226.27	76.75	153.29
Nb	170.07	265.20	275.54	80.82	151.57
Mo	1.19	2.11	2.68	1.01	1.74
Cs	0.72	0.31	. 0.27	0.99	0.39
Ba	695.44	2386.92	1772.93	207.00	842.58
Hf	3.01	4.67	4.78	1.84	3.41
Ta	9.73	10.57	11.09	4.81	6.73
Pb	8.47	9.38	2.92	8.91	4.48
Th	12.39	25.40	26.22	5.54	15.03
U	2.81	5.01	3.96	1.46	2.60
La	102.22	217.21	214.64	41.84	114.08
Ce	174.50	375.52	374.21	78.47	200.69
Pr	18.11	38.54	39.03	8.52	21.06
Nd	62.19	129.64	131.63	30.34	73.21
Sm	8.36	16.53	17.21	4.41	9.83
Eu	2.10	4.21	4.20	1.14	2.52
Gd	4.22	7.68	8.35	2.61	5.04
Tb	0.59	1.14	1.19	0.38	0.73
Dy	2.59	4.66	4.80	1.73	3.06
Ho	0.39	0.67	0.70	0.29	0.46
Er	0.84	1.34	1.39	0.68	0.92
Tm	0.11	0.17	0.17	0.10	0.12
Yb	0.58	0.88	0.90	0.57	0.66
Lu	0.08	0.12	0.12	0.09	0.09
87Sr/86Sr	0.705222	0.705041	0.705358	0.705264	0.705394
143 Nd/ 144 Nd	0.512576	0.512593	0.512615	0.512583	0.512588
²⁰⁶ Pb/ ²⁰⁴ Pb	20.325	20.445	24.744	19.238	26.211
²⁰⁷ Pb/ ²⁰⁴ Pb	15.592	15.594	15.922	15.590	15.961
²⁰⁸ Pb/ ²⁰⁴ Pb	40.577	41.189	51.016	38.957	50.351

Locations: UV8M - Udachnaya, A3 and A4 - Aikhal, M3 - Mir and

I-1 - Internationalaya