

Re-Os and Sm-Nd Isotope Systematics of Alkaline Ultramafic Rocks, Xenoliths and Macrocrysts from the Earahedy Basin, Yilgarn Craton

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Introduction

In order to understand better the role of the continental lithospheric mantle and incompatible element enrichment processes in the generation of alkaline ultramafic rocks we have obtained Carius tube/N-TIMS Re-Os isotopic data from kimberlites, melnoites, oxide macrocrysts and lherzolite xenoliths from the Earahedy Basin, Western Australia. We have previously demonstrated that marginal, on-craton Yilgarn melnoites and picroilmenite macrocrysts have a source prehistory coincident with metasomatism-accretion of fertile lithospheric mantle during Proterozoic orogenesis (Graham et al., 1996). Re-Os isotopic data from ilmenite megacrysts and pyroxenite xenoliths from South African kimberlites (Olive et al., 1997) support our previous conclusions that younger lithospheric mantle can be accreted to the continental lithospheric mantle by upper mantle incompatible element enriched melts.

Background Geology

The northern margin of the Eastern Goldfields Province is defined by rock sequences of the Earahedy Basin (Fig. 1). The basin is one of the structural features which belong to the 2200 - 1600 Ma Capricorn Orogen. Basin basement rocks are probably Archaean Yilgarn granitoid-greenstones and the internal structure of the Capricorn Orogen is interpreted to reflect Pilbara-Yilgarn collision which resulted in basin development (Tyler et al. in press). Three alkaline ultramafic provinces of different ages occur within and near the Earahedy Basin. Emplacement ages range from ~1900 - 1700 Ma (Nabberu kimberlites: Shee et al., 1996), 1329 ± 9 Ma (Jewell kimberlites: unpubl. Rio Tinto K-Ar phlogopite data), to 305 Ma (Bulljah melnoites: Hamilton and Rock, 1990, SHRIMP zircon).

Geochemical and Isotopic Results

Whole rock kimberlite and melnoite samples are enriched in LILE, HFSE and LREE, consistent with literature data for alkaline ultramafic rocks. The Nabberu kimberlite has a flat primitive mantle normalised PGE pattern, similar to continental lithospheric mantle peridotites, kimberlites and orangites (Fig. 2). The Bulljah melnoite has a fractionated PGE pattern, similar to patterns observed in mantle melts. The negative Pd anomaly, a feature observed in all melnoites from the Yilgarn Craton, may result from fractionation of spinel or be a characteristic of Yilgarn melnoite sources. Re/Os concentration ratios for kimberlites are variable and range from 0.035 to 0.138, while Re/Os ratios for the melnoites are significantly higher (0.107 to 1.261) indicating a higher degree of partial melting or smaller percentage of xenocrysts than the kimberlites (Fig. 3).

Sm-Nd isotopic data from the Nabberu kimberlites and Bulljah melnoite provide evidence for two distinct sources. The Nabberu kimberlites have an unradiogenic initial ϵ_{Nd} isotopic composition ($\epsilon_{Nd} = -7$). The Bulljah melnoite also has an unradiogenic initial isotopic composition ($\epsilon_{Nd} = -20$), but is radiogenic at 1900 Ma. Sm-Nd T_{DM} model ages for the two rock types also show distinct differences, the Nabberu kimberlites have model ages of ~2700 Ma while the melnoite has model age of 2100 Ma. Sm-Nd isotopic and PGE data argue strongly against a genetic link between the two rock types, despite their incompatible trace element affinities and close spatial association. Although the ~2700 T_{DM} model ages from the Nabberu kimberlite are similar to Marymia Dome crustal ages contamination is discounted as alkaline ultramafic rocks have Sm and Nd concentrations greater than average crust. The data may therefore be representative of anomalous enriched regions of continental lithospheric mantle beneath this region.

Re-Os isotopic data for the three provinces reveal an interesting and significant evolutionary history for the source regions of these alkaline ultramafic rocks. Whole rock data for kimberlite and their mantle xenoliths are isochronous (Fig. 4). The chondritic initial $^{187}\text{Os}/^{188}\text{Os}$ isotopic

ratio of 0.1159 and the 2061 ± 248 Ma age are consistent with the active margin phase of Pilbara and Yilgarn Craton accretion (Tyler and Thorne, 1990). We interpret this line to represent a "mantle isochron", because the isotopic composition of kimberlites can be dominated by disaggregated lithospheric mantle (Pearson et. al., 1995). Thus, this isochron reveals the age of the shallow mantle beneath the Earaaheedy Basin.

Re-Os isotopic data for whole rocks and spinels from the Bulljah melnoite yield an isochron with an age of 1281 ± 115 Ma, MSWD = 13 (Fig. 5). We interpret this age as the age of emplacement because the 305 Ma emplacement age was determined from 3 zircons which have U contents far greater ($U > 100$ ppm) than those typical alkaline ultramafic rocks, also the Re-Os age is within error of the K-Ar phlogopite emplacement age of the Jewill kimberlite. The sub-chondritic initial $^{187}\text{Os}/^{188}\text{Os}$ for the isochron is inconsistent with the near chondritic and radiogenic Os isotopic ratios observed in asthenospheric and plume mantle melts respectively. The low isochron initial is however consistent with the melnoites being the melt of continental lithospheric mantle that had possessed a lower Re/Os ratio than the asthenospheric mantle for a long period of time. The low isochron initial $^{187}\text{Os}/^{188}\text{Os}$ will allow a T_{RD} model age to be calculated. This model age will have a geological meaning only if it approximates the lithospheric mantle source of the melnoite and if the source was devoid of Re at 1281 Ma. The initial $^{187}\text{Os}/^{188}\text{Os}$ isotopic composition has a calculated Re-Os T_{RD} minimum Re depletion model age of ~ 2200 Ma, within error of the kimberlite isochron (Fig. 6). The melnoite isochron therefore places important constraints on the nature of mantle sources for alkaline ultramafic rocks from the Earaaheedy Basin.

The lack of diamond from the Bulljah melnoite is consistent with a larger melt component in these rocks compared to diamondiferous kimberlites of the same age located further west. The Re-Os data, in contrast to the Sm-Nd and PGE data, confirm a genetic link between kimberlites and melnoites. Moreover, the data show that it is highly likely that both rock types are derived from the lithospheric mantle and that source melting must be achieved through processes other than simple plume-continental lithospheric mantle interaction.

Conclusions

Re-Os and Sm-Nd isotopic data for alkaline ultramafic rocks from the Earaaheedy Basin are consistent with accretion of new lithospheric mantle during orogenesis ($\sim 2200 - 2000$ Ma). Furthermore, emplacement ages coincide with upper crustal basin extension (Nabberu ~ 1700 Ma; Bangemall 1400 - 1300 Ma). Archaean lithospheric mantle of the proto-Earaaheedy Basin may have been extensively thinned and replaced by accreted melt depleted oceanic lithospheric mantle at ~ 2100 Ma. Incompatible element enrichment of the sources also occurred at this time. Re-Os isotopic data from the Bulljah melnoite suggest a genetic relationship between the kimberlites and melnoite. No diamonds are found in association with the melnoite because of the larger degree of partial melting.

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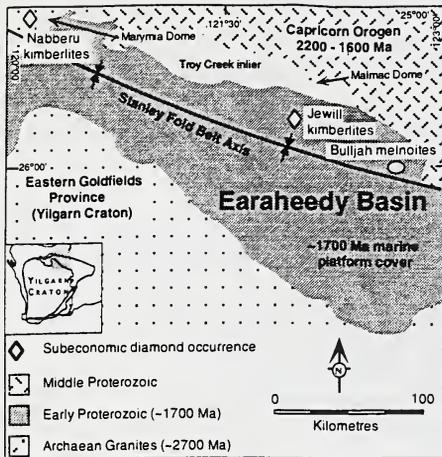


Figure 1. Eoraheedy Basin Geology

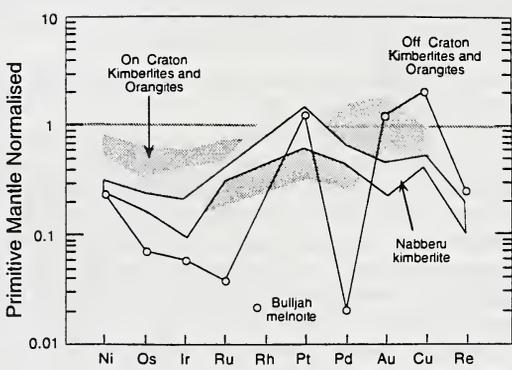


Figure 2. PGE abundances

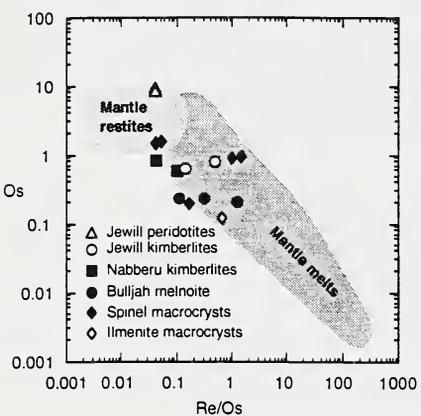


Figure 3. Re-Os mantle array

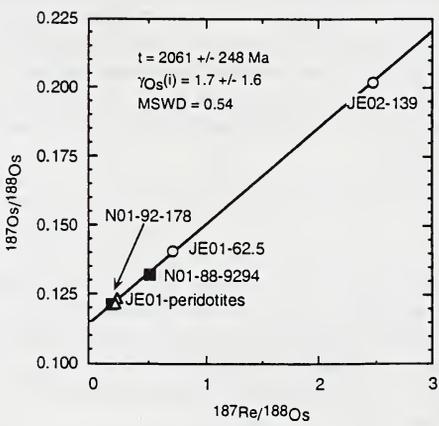


Figure 4. Kimberlite isochron

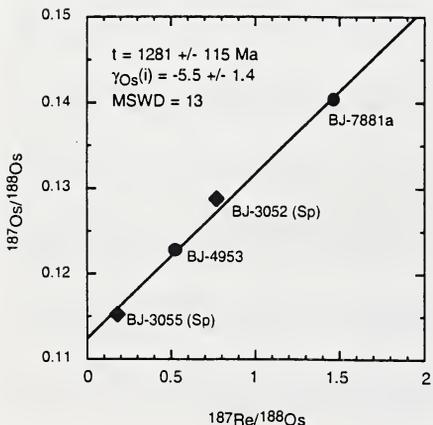


Figure 5. Melnoite isochron

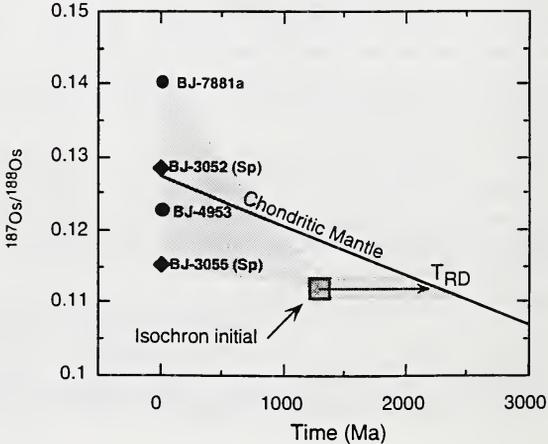


Figure 6. Bulljah Os evolution