ECLOGITE XENOLITH WITH EXSOLVED SANIDINE FROM THE PROTEROZOIC KURUMAN KIMBERLITE PROVINCE, NORTHERN CAPE, R.S.A.

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The zero kimberlite is one of twelve known occurences of kimberlites and related rocks of the Kuruman Province, northwest Cape, South Africa. Radiometric dating has demonstrated that these bodies are 1600 to 1700 Ma in age (Shee et al., 1989), and are thus amongst the oldest known kimberlites. They occur as small eroded pipes or dykes located on the western edge of the Kaapvaal Craton (greater than 2500 Ma) adjacent to the younger Kheis belt (about 1800 Ma) of the northern Cape Province. Zero is an oval-shaped pipe, 300m by 200m, intrusive into the early Proterozoic Campbell Rand dolomite of the Transvaal Super-group. Texturally the kimberlite is a macrocrystic hypabyssal type, and is mineralogically classified as an opaque mineral-rich phlogopite calcite group I kimberlite.

The kimberlite is notable for the high mantle-derived xenolith content, but the pipe is capped by extensive calcrete, and preserved material is available only from drill core. The xenolith suite comprises equal proportions of eclogite and peridotite, the latter rocks having been studied by Shee et al. (1989) who documented evidence for a hot, perturbed geotherm. The eclogite nodules have not been studied previously in any detail. For this study a single eclogite xenolith was extracted from the Zero drill core for petrologic study, the sample being one of two initially chosen for their uncommonly large size, unusual degree of freshness, and unusual textural relations between garnet and clinopyroxene. Careful petrographic study of the speciman indicates that it is essentially a large clinopyroxene megacryst that has exsolved garnet and sanidine. Sanidine has only rarely been described as occuring in mantle eclogite (Smyth and Hatton, 1977; Smith, 1977; Ater et al., 1984), and this is the first known case where sanidine is a clear exsolution product, occuring both as discrete grains at clinopyroxene and garnet grain boundaries and as crystallographically controlled exsolutions blebs within clinopyroxene. Prior to exsolution the megacryst must have contained substantial potassium in solid solution, implying a very high pressure origin.

Average major elements analysis and standard deviations of the mineral phases (Table 1) demonstrate that they are relatively homogeneous. Garnet is relatively magnesian and calcium-poor, and in this regard is distinctive compared to previously documented sanidinebearing samples from kimberlite. The example described by Smyth and Hatton (1977) and Wohletz and Smyth (1984) is a grospydite with considerably greater Ca. Of twelve samples noted by Ater et al. (1984) from the Colorado-Wyoming kimberlites, nine are kyanite- and/or corundum-bearing and the other three are significantly more Fe- and Ca-rich than the Kuruman example. Clinopyroxene in the Kuruman sample has low to moderate Na₂O of 1.7%. Na₂O in garnet (0.05%) and K₂O in clinopyroxene (0.03%) are low compared to type 1 diamond-bearing assemblages (McCandless and Gurney, 1989), and Cr₂O₃ is relatively high in both minerals. All of these chemical features indicate a type 2 association for the sample, as is the case for most other documented sanidine-bearing eclogites. The K_b value is 3.24, and

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for an assumed pressure of 50 Kb the equilibration temperature is 1000⁰C. This falls slightly below the peridotitic geotherm calculated by Shee et al. (1989).

Clinopyroxene and garnet are both enriched in Ce (20.2 and 9.5 ppm, respectively), and relative to La and Nd the clinopyroxene has a distinct negative Ce anomaly with overall LREE enrichment (La/Eu = 4.54). High Ce in garnet also gives a LREE enriched pattern, with HREE enrichment from Gd to Lu. Further isotopic and trace elements analysis are in progress.

Given the setting and age of the Kuruman kimberlites, data from this xenoliths and as yet unstudied additional members of the eclogite assemblage could have important bearing on the understanding of the origins and evolution of lithospheric mantle beneath the Kheis belt and marginal to the craton for which various subduction-type models have been postulated.

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Table 1. Electron Microprobe Analysis of Garnet, Clinopyroxene and standard deviations for N analysis.

	GAR	CPX	SAN
N	6	10	3
SiO ₂	40.9 +/3	54.8 +/03	62.2 +/3
TiO ₂	0.06 +/02	0.10 +/02	-
Al ₂ O ₃	22.8 +/2	2.94 +/09	18.4 +/3
Cr ₂ O ₃	0.25 +/03	0.15 +/02	-
FeO	16.2 +/1	4.88 +/12	0.77 +/05
MnO	0.45 +/02	0.07 +/02	-
MgO	14.9 +/2	14.5 +/2	-
CaO	4.59 +/12	20.3 +/1	-
Na ₂ O	0.05 +/02	1.71 +/06	0.37 +/10
K ₂ O		0.03 +/01	15.6 +/3