

CARBON ISOTOPE COMPOSITION OF GRAPHITE IN MANTLE ECLOGITES.

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The carbon isotopic composition of diamond is reasonably well-known from measurements on hundreds of diamonds of known paragenesis. With very few exceptions, $\delta^{13}\text{C}$ values of peridotite-suite diamonds are in the restricted range -2 to -9‰_{PDB}, whereas $\delta^{13}\text{C}$ for eclogitic diamonds ranges from approximately +2 to -34‰_{PDB} (e.g., Sobolev et al., 1979; Deines, 1980; Harris, 1987). In contrast to this large data set, few isotopic data have been published for either eclogitic or peridotitic mantle-derived graphite. Values of $\delta^{13}\text{C}$ for primary graphites from peridotite xenoliths (-4.8 to -10.04‰; Kropotova and Fedorenko, 1970; Pearson et al., 1990; Schulze and Valley, in press) are similar to those from peridotitic diamonds. Existing carbon isotope data for graphite in mantle eclogites from Mir (Kropotova and Fedorenko, 1970), Roberts Victor (Deines et al., 1987), and Orapa (Deines et al., 1991) are, with one exception, in the "typical mantle range" (-3.98 to -8.7‰_{PDB}). Graphite from a Schaffer, Wyoming kyanite eclogite (Schulze and Valley, in press) is unusually light, with $\delta^{13}\text{C} = -14.31\text{‰}$. An additional anomalous graphite analysis ($\delta^{13}\text{C}$ of -20.3‰) was reported by Deines et al. (1991) for a corundum-garnet-spinel assemblage of uncertain affinity, also from Orapa. To enlarge the data base for mantle-derived graphites, we present carbon isotope data for graphites from 23 eclogite xenoliths, from the Bellsbank and Jagersfontein kimberlites in South Africa and the Orapa and Letlhakane kimberlites in Botswana.

Most of our new data (Table 1) are normally distributed about a $\delta^{13}\text{C}$ value near -5 to -6‰. Interesting features in this new data set include a very light value of -12.50‰ for an eclogite from Letlhakane, and the fact that 3 of the 8 Bellsbank samples have a $\delta^{13}\text{C}$ value near -2.9‰. These latter values are unusually heavy, relative to the strong peaks for both eclogitic and peridotitic diamonds at approximately -6‰, but are similar to many of the $\delta^{13}\text{C}$ values for loose diamonds of unknown affinity from Bellsbank. One third of the diamonds from the Dan Carl Mine at Bellsbank reported by Deines (1980) have $\delta^{13}\text{C}$ values in the range -2 to -3‰.

Graphites from mantle eclogites thus appear to have an overall $\delta^{13}\text{C}$ distribution similar to eclogite-suite diamonds. In both suites there is a strong peak near -6‰, with a significant number of samples at lower and higher values, although the relatively few graphites do not define as wide a range of $\delta^{13}\text{C}$ values as does the larger eclogitic diamond population.

The origin of the isotopically light and heavy eclogitic diamonds is controversial. Considered in isolation, the range of $\delta^{13}\text{C}$ values for both eclogitic diamonds and graphites could be explained by several different processes. These include precipitation of native carbon from CO_2 and CH_4 vapours through Rayleigh fractionation, isotopic inhomogeneity in the mantle remaining since the accretion of the Earth, and subduction

of oceanic basalt containing crustal biogenic and carbonate carbon (e.g., Deines, 1980; Kirkley and Gurney, 1989). Only the subduction hypothesis, however, adequately accounts for many of the other characteristics of mantle eclogites, characteristics such as positive Eu anomalies and anomalously high and low $\delta^{18}\text{O}$ and $\delta^{34}\text{S}$ values, that defy explanation in terms of mantle igneous processes.

We suggest that the anomalous $\delta^{13}\text{C}$ values for graphite in mantle eclogites support the subduction hypothesis. Furthermore, the model of Helmstaedt and Schulze (1989) of the formation of the subcratonic lithosphere by tectonic emplacement of ocean-floor material beneath the cratons can be extended to include even the shallow portion of the lithosphere represented by rocks from the graphite stability field.

Table 1. Carbon isotope composition of graphite from mantle eclogites, analyzed at the University of Wisconsin.

Location	Sample	$\delta^{13}\text{C}_{\text{PDB}}$
Letlhakane	JSL-200	-5.54
	JSL-201	-5.69
	JSL-202	-5.37
	JSL-204	-7.68
	JSL-206	-12.50
	JSL-207	-5.68, -5.75
	JSL-208	-5.98
Jagersfontein	K7/555	-5.90
Bellsbank	K64/124	-2.94
	K64/128	-5.69
	K64/129	-7.04
	K64/133	-2.95
	K64/134	-2.84
	K64/147	-5.38
	K64/149	-4.96
	K64/169	-5.47
Orapa	K1/399	-6.55
	K1/400	-6.19
	K1/401	-6.23
	K1/402	-4.64, -4.71
	K1/403	-5.67
	K1/404	-6.29
	K1/406	-6.00

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