THE GEOLOGY AND MINERALOGY OF SOME KIMBERLITES IN THE MWADUI AREA

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The Mwadui kimberlite-bearing area is located in the Tanzania Archean Craton (2.5 - 3.0 Ma). The craton contains the oldest rocks known in Tanzania namely greenstones (basic and acid metavolcanics, banded ironstones and metasediments) and granitoids. The majority of the kimberlites lie within the craton and the Mwadui area is one with the highest number of fertile kimberlites. Out of the 67 known kimberlites in the latter area, 25 are diamondiferous, and only 3 have been economically exploited for diamonds. Among the three economic kimberlites is the Mwadui kimberlite pipe covering an area of about 146 ha making it the largest kimberlite in the world. Since its discovery (1940) the Mwadui pipe has produced 3.4 tons of diamonds valued at US\$1200 million.

The Mwadui area kimberlites were emplaced into the granitoids and greenstones during the Early Tertiary times (40-53 Ma; Davis, 1977; Raber, 1978; Shee, 1981). A considerably older age of Mid-Proterozoic (1097 Ma) has been obtained for one kimberlite in the Bubiki area (Bristow, 1986).

A preferred structural trend of the kimberlites is NE-SW which coincides with the trend of dolerite dyke swarms of Upper Proterozoic age. The Mwadui pipe lies at the intersection of the NE-trending dolerite dykes and a NWstriking shear zone. Close to the pipe is a major NNE-trending lineament. A structural trend between N and NNE appears to be associated with major kimberlitic intrusions e.g. Mwadui, Mhunze and Nyamigunga.

Macroscopic and microscopic study done recently has permitted in the mineralogical and textural classification of kimberlites in the Mwadui area. The mineralogical classification follows Skinner and Clement (1979) and is based on the predominant macrocrysts whereas the textural classification (Clement and Skinner, 1979) gives the facies (crater, diatreme and hypabyssal).

The study of the geology of the Mwadui pipe has shown that all the facies of a true kimberlite are preserved namely crater, diatreme, and hypabyssal (Gobba, 1989). In the crater facies are found both epiclastic and pyroclastic kimberlitic sediments (tuffs, shales, mudstones). The crater facies rocks are overlain by superficial deposits (gravels, silcretes calcretes, and clays). Below the bedded crater facies rocks are the diatreme facies tuffisitic kimberlite breccias which are characterized by abundant unaltered xenoliths, discrete and fragmented autoliths of hypabyssal kimberlite and xenocrysts from from deeper crustal/mantle sources. True hypabyssal facies within the main pipe has not been exposed. However, there are kimberlite dykes to the N and S of the main pipe which show mineralogical and textural properties of hypabyssal kimberlite. The dyke to the north has a NE-SW orientation and is an altered serpentine-diopside kimberlite. The southern dyke has a SE trend and is an altered kimberlite containing serpentine and clay minerals.

A few other kimberlites show all the three facies (e.g. Mhunze and Nyamigunga) whilst most of the remaining are eroded to deeper levels removing some of the facies (mainly the crater facies).

Many Mwadui area kimberlitic intrusions are characterised by an absence of macrocrystal olivine and it is suggested that these could be related rocks here after called parakimberlites, transitional between kimberlites and other ultramafic rocks (Gobba and Saxby, 1982; Gobba and Edwards, 1983). The criteria used in this study to classify a rock as a para-kimberlite is the morphology of olivine grains and the presence of groundmass minerals such as nepheline or analcime (Shee, 1980). Some kimberlitoids have been found in Mwadui area e.q. 65K34 kimberlite pipe. In this intrusion extensive mineral alteration to serpentine, carbonates and clay minerals are common. Pseudomorphs of olivine and monticellite are found, spinels and apatite are ubiquitous. At Bubiki (48K5 kimberlite pipe) olivines have unusual shapes and perovskites are unusually large. Overall the rock is considered to be a marginal kimberlite possibly transitional to an olivine melilite type rock association.

Some of the occurrences may represent aphanitic kimberlites and may have come about through filter pressing process. These however are not common and usually occur in specific geological settings such as dykes and sills. The fairly widespread occurrence with this texture is unusual for kimberlites. Several possibilities are given by Wyatt (1983) who suggested that they could be of relevance if these rocks are kimberlites. Firstly, they could have unusual structural control which could account for the removal of coarse constituents. Alternatively the absence of macrocrysts could be an intrinsic feature of the Mwadui area as a whole. It is therefore interesting to speculate that since the Mwadui kimberlite is almost completely preserved, we may be observing, in the Mwadui area, the terminal features of kimberlitic episodes that may not normally be observed in other areas. However considering combined evidence to date it seems probable that these occurrences are related rocks rather than aphanitic kimberlites. In regard to this it would seem that the presence of Mwadui true kimberlite in a province with fairly numerous occurrences of kimberlite-related rocks is unusual.

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