

D Xenoliths and geometers

D1 ENSTATITE-DIOPSIDE SOLVUS TO 60 KB

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The width of the pyroxene solvus in the simple system enstatite-diopside as a function of pressure and temperature is an important constraint for the use of the composition of natural pyroxenes from lherzolites as a geothermometer. Experimental data have been obtained so far only up to 40 kb (only at 1200°C at this pressure). Thermodynamic models based on high pressure experiments and calorimetric data should not be extrapolated above 40 kb according to the authors (e.g. Lindsley, 1980). However pressures up to 90 kb have been estimated in some garnet lherzolites so that an extrapolation of the pyroxene solvus above 40 kb is necessary to estimate the temperature of many garnet lherzolites. With a well-calibrated belt-apparatus (against the silver-melting curve and the quartz-coesite transition) I have now embarked to determine the pyroxene solvus up to 60 kb. Pressures are controlled within 300 bars and the temperature to $\pm 1^\circ\text{C}$ (estimated accuracy ± 0.5 kb and $\pm 10^\circ\text{C}$). Starting materials are mixtures of synthetic enstatite and diopside and for reversals glass of the same bulk composition. So far experiments have been carried out at 40 kb at 1100°C, 1300°C and 1500°C and at 1300°C at 50 kb and 60 kb. The compositional brackets from the experiments for the solvus at 1300°C are given in Table 1. The data are consistent with experimental

	$x_{\text{CaMgSi}_2\text{O}_6}^{\text{En}}$	$x_{\text{CaMgSi}_2\text{O}_6}^{\text{Di}}$	data given in the literature.
40 kb	0.052	0.78-0.8	They indicate
50 kb	0.045-0.048	0.785-0.815	that, at 1300°C,
60 kb	0.036-0.038	0.79-0.83	the influence of pressure on

the position of the solvus becomes less with increasing pressure. An extrapolation of the currently available models on the P, T-dependence of the pyroxene solvus shows that temperature estimates can be too high by up to 250°C at 60 kb.

Lindsley, D.H., *Min.Soc.Rev.Min.* 7 (1980) 289

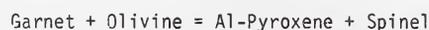
D2 GEOTHERMOMETERS AND GEOBAROMETERS FOR SPINEL-BEARING ULTRAMAFIC ROCKS

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A series of two pyroxene, spinel-bearing assemblages have been synthesized over the

temperature and pressure range from 900°C to 1300°C and 10Kb to 15Kb, respectively. Systematic compositional variations have been introduced by varying the composition of the reactants in the following synthesizing reaction.



In this way it has been possible to experimentally examine the competing effects of Cr and Al substitution in coexisting pyroxenes and spinels.

The synthesized products have been analyzed with an electron microprobe and used to empirically define expressions that describe exchange reactions between coexisting phases. The results yield a series of geothermometers for Al/Cr exchange between coexisting pyroxenes and spinels, Fe/Mg exchange between coexisting silicates and silicates and spinels. For the compositions used observed pressure dependence of the orthopyroxene-clinopyroxene solvus yields a useful geobarometer for these phase assemblages

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D3 THE OLIVINE-CLINOPYROXENE GEOBAROMETER: IMPLICATION FOR THE THERMAL STRUCTURE OF THE UPPER MANTLE NEAR THE RIO GRANDE RIFT

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Ca-Mg exchange between olivine and clinopyroxene has been studied in the system MgO-CaO-SiO₂ in the temperature range 1100-1500°C and pressure range 9-41 kb. Ca in olivine decreases with increasing pressure as a result of Ca going into the higher coordination M2 site of clinopyroxene. For the Ca-Mg exchange reaction between olivine and clinopyroxene, $\Delta V^\circ = 0.249$ j/bar/mole, sufficient for pressure estimates precise to ± 3 kb if temperatures of equilibration are independently known.

The olivine-clinopyroxene geobarometer can be used as an independent check on results from pyroxene barometry applied to garnet lherzolites. Also, it does not encounter the problems inherent in pyroxene barometry of spinel lherzolites. The olivine-clinopyroxene barometer has been applied to

xenoliths from several localities to determine local geotherms, after making assumptions about the activities of various components in the natural solid solutions. Results on southern African peridotites compare favorably with pressure estimates made by independent geobarometers and with calculated continental conduction geotherms. Results from southwestern United States, however, indicate that the geotherm is elevated, though not to the extent of mid-ocean ridge convection geotherms. This may be due to an intermediate thermal state, involving both convection and conduction, resulting from the slow rate of spreading occurring along the Rio Grande Rift.

D4

GARNET LHERZOLITES FROM THE HANAUS-I AND LOUWRENSIA KIMBERLITES, NAMIBIA

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The Gibeon cluster of Namibian kimberlites is emplaced into the Orange River Belt which has accreted to the Transvaal craton. These "off-craton" kimberlites are non-diamondiferous and are younger than the diamond bearing "on-craton" kimberlites. The Hanaus-I and Louwrensia kimberlites each contain a bimodal suite of upper mantle-derived garnet lherzolite xenoliths characterized by a coarse granular or a porphyroclastic-mosaic texture. The Louwrensia pipe in addition contains garnet harzburgites in which the orthopyroxenes have exsolved lamellar clinopyroxenes. Deformed lherzolites are not iron-enriched relative to the deformation-free types. Temperatures and pressures of equilibration calculated by the Wells-Wood method are from 915-1050°C at 27-36 kb and 915-1010°C at 29-41 kb for coarse lherzolites from Hanaus and Louwrensia respectively, and from 1000-1225°C at 26-39 kb and 1010-1125°C at 33-34 kb for porphyroclastic-mosaic types from Hanaus and Louwrensia respectively. The coarse types from both localities have similar equilibration P and T's to coarse lherzolites from on-craton kimberlites and plot on the lower non-inflected limb of the South African continental shield "paleogeotherm" defined by such lherzolites. The deformed lherzolites plot randomly above this geotherm and do not define an inflected limb. These higher equilibration temperatures are considered to be due to heating, shearing and recrystallization of mantle during kimberlite formation. The lherzolite mineral compositions and equilibration conditions indicate that the Namibian kimberlites have been derived from similar depths to on-craton kimberlites and have incorporated similar mantle material. The lack of diamonds is not a consequence of derivation from atypical mantle or from regions with an unusual geothermal gradient.

D5

GARNET PERIDOTITES FROM WILLIAMS KIMBERLITES, NORTH-CENTRAL MONTANA, USA

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Two Williams kimberlites, 250x350m and 37x390m, in the eastern part of a swarm of 30 middle Eo-

cene alnoitic diatremes, contain xenoliths of garnet-bearing lherzolites, harzburgites and dunites. Temperatures were calculated by the Lindsley-Dixon 20 kb method for lherzolites and by the O'Neill-Wood method for harzburgites and dunites and pressures were calculated by the MacGregor method, or were assumed to be 50 kb for dunites. Most peridotites equilibrated at 1220-1350°C and 50-60 kb, well above a 44 mW/m² shield geotherm and on or at higher P than the graphite-diamond boundary. Three lherzolites show low T-P (830-940°C, 23-42 kb) and are close to the shield geotherm. All three low T-P lherzolites have coarse texture whereas the high T-P cluster has both coarse and porphyroclastic textures, indicating a range of conditions of deformation and recrystallization in a restricted high T-P range. Maximum size of large strained olivines is as much as 2 cm. The tiny size (0.1-0.2 mm) of granulated and euhedral olivines in several xenoliths shows that deformation was occurring just prior to incorporation in kimberlite and that ascent was rapid enough to retard further coarsening of fine-grained olivine. For other high T-P peridotites, cessation of deformation before inclusion in kimberlite is suggested by larger (2mm) euhedral olivines in a matrix of fine granulated olivine or by optical continuity of large and nearby small olivines. Two low T-P lherzolites contain 5-8mm clots of moderate-Cr garnet + Cr-spinel + Cr-diopside inferred to form by reaction of an initial high-Cr garnet brought into the garnet-spinel stability field. This suggests a slower ascent or pause in ascent, compared with other peridotites containing inclusion-free, high-Cr garnets. Textural and compositional variations of peridotites are compatible with kimberlite generation and ascent during dynamic diapiric perturbation of the upper mantle.

D6

MANTLE GARNET-SPINEL TRANSITION ZONE DEMONSTRATED BY XENOLITHS FROM COLORADO-WYOMING KIMBERLITES

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Two groups of xenoliths containing coexisting garnet and spinel have been recovered from Colorado-Wyoming kimberlites. The first group consists of pyroxenites with coarse-grained, pale green, aluminous spinel that is relatively Mg-rich and Cr-poor. Garnet in these xenoliths appears to be exsolved from aluminous pyroxene, and occurs as lamellae and as grain boundary rims surrounding pyroxenes and spinels. Xenoliths of the second group are peridotites containing fine-grained vermicular, reddish-brown spinel that is more enriched in Fe and Cr relative to Al. Some garnet in these nodules also occurs as exsolution lamellae in pyroxene, but most engulfs spinel and probably formed by the reaction *orthopyroxene + clinopyroxene + spinel = olivine + garnet* as suggested by MacGregor (1970). Utilizing MacGregor's (1974) orthopyroxene geobarometer, mineral equilibration depths of 50-100 km and 25-60 km have been calculated for the peridotite and pyroxenite xenoliths, respectively. Temperature estimates range between 590-800°C. Garnet in both xenolith groups apparently formed during or immediately prior to kimberlite emplacement and the