PALEOMAGNETISM AND THE AGES OF KIMBERLITES EXEMPLIFIED BY THE FOUR PIPES OF YAKUTIA

Zhitkov<sup>(1)</sup> A.N., Savrasov<sup>(2)</sup> D.I.

(1) VostSibNIIGGiMS, Geological Committee of Russia, Irkutsk, Russia; (2) Botuobinsky expedition, Mirny city, Russia

The oriented samples were collected in four Middle Paleozoic economic diamond-bearing bodies of Malya Botuoba (Sputnik, named after 23rd Congress) and Alakit-Markhinsky (Sytykanovsky, Yubileiny) regions. The collections represent all main varieties of kimberlites of every pipe (661 samples). The magnetometric measurements were made by highprecision spin-magnetometer JR-4 and kappabridge KLY-2. Experiments on thermal demagnetization and demagnetization by alternating magnetic field were conducted in the facility of magnetic vacuum with a precision of the field compensation + 10 nT. A step of scalariform demagnetization was 25-500 C and 20-40 kA/m, accordingly. The diagnostics of the magnetism bearers was carried out on the curves of normal

Table 1

Magnetic properties of kimberlites in natural occurrence

	М	б	Ml	3	J	D	К	A <sub>95</sub>	ρ <sub>зу</sub>
I In H Q	465 578 377 2.7	Pip 95 92 28 0.2	"After 225 293 241 2.4	23rd ( 1.11 1.12 1.10 1.07	Congress -31.4 -71.2 -	5", N-13 5.7 22.4 -	38   1.9   3.9   -   -	12 7 -	42 29 -
I In H Q	401 396 280 1.9	60 64 33 0.2	78 36 74 0.8	1.15 1.22 1.14 1.12	$\begin{vmatrix} 111 \\ 51.8 \\ 3.9 \\ - \\ - \\ - \\ - \\ - \\ - \\ - \\ - \\ - \\ $	$  354.6 \\ 347.5 \\ - \\ - \\ - \\ - \\ - \\ - \\ - \\ - \\ - \\ $	3.1 1.8 -	7 12 -	33 44 -
I In H Q	2092 781 2465 1.42	184 86 205 0.17	786 285 871 0.54	1.11 1.12 1.11 1.09		$\begin{vmatrix} & N - 2 \\ & 344 \\ & 340 \\ & - \\ $	± 3 2 	6 10 - -	34 44 -
I In æ Q	2582 618 3337 0.27	88 48 82 0.02	2297 372 3084 0.20	1.05 1.10 1.04 1.08	80 80 70 - -	$ \begin{vmatrix} 1 \\ 1 \\ 11 \\ - \\ - \\ - \end{vmatrix} $		1   8   -   -	5 31 -

I - ummary magnetization, In - natural residual magnetization, pprox - magnetic susceptibility, Q - Kenigsberger coefficient, M - average arithmetic, M1 - average geometric, 6 - standard deviation,  $\varepsilon$  - standard multiplier of lognormal distribution, J(D) - inclination (declination) of the single vector resultant, K - clustering, A<sub>95</sub> - radius of confidence circle for 95% probability (after R.Fischer, 1957),  $\rho_{39}$  - radius of a single circle of dispersion (spherical standard). Values I, In are given in mA/m, pprox - in 10<sup>-5</sup> CI units; J, D, A<sub>95</sub>,  $\rho_{39}$  - degrees; Q,  $\varepsilon$ , K - dimensionless. 695 magnetization in the fields up to 600 kA/m. Quantitative
paleomagnetic dating was preformed via comparison with the
trajectory of apparent migration of the Siberian platform.
 A complete spectrum of magnetic parameters was obtained and
statistically processed. The values and dispersion of magnetic
properties of kimberlite in the natural occurrence are provided
in Table 1. All studied bodies were found to contain the rock
varieties retaining primary natural residual magnetization which
can be easily interpreted. An instance of thermal demagnetization
of kimberlite with a dominant primary component of natural
remanent magnetization is given in the Figure.



Figure. Results of thermal demagnetization. The porphyry kimberlite of endocontact zone of "Sytykan" pipe.

A particular feature of the kimberlites studied is a wide spectrum of limiting temperatures and Curie points bearing natural remanent magnetization (from 250 to 6500 C). This indicates presence in source rocks of hypogenous magnetic minerals of different age (magnetite, hemoilmenites with different ratios of Fe-Ti components).

The results of paleomagnetic studies and quantitative polar paleomagnetic dates are listed in Table 2. The conditional paleomagnetic directions and poles were obtained for four pipes. The primary origin of natural residual magnetization is substantiated by a set of criteria including the most reliable method of comparison of paleomagnetic directions obtained for kimberlites and host sedimentary rocks in the contact zone (test of burning). It was revealed that remagnetization of hosting rocks by kimberlite occurs only in the vicinity from the contact (the first centimeters). A significant and unexpected result was inverse polarity and the primary natural residual magnetization of all studied varieties of the Middle Paleozoic kimberlites of the two separated kimberlite fields. There is every ground to believe that this feature can be bound or dominant in other bodies of the same or close age. A stable inverse polarity of geomagnetic field contributes to origination of the Middle Paleozoic kimberlites.

Table 2

Paleomagnetic directions, poles and dates of Yakutian kimberlites

		Pa d:	leoma irect	agneti tions	С	Region		Paleomagnetic poles				T mln.
	N/n	J	D	k	α95	φ	λ	ф	Λ	A1	A2	y15.
1'. 2.	16/49 15/45	-70  -66	Ma 318 326 Alaki	laya B 121.1 39.5 It-Mar	otuol 3.2 5.8 khins	ba regi 62.5 62.5 sky reg	ion 113.5 113.5 gion	32 24	141 138	5.5 9.4	4.8 7.6	345±9 348±11
3. 4.	14/60 11/37	-64 -63	322 324	42.0 100.2	$5.8 \\ 4.2$	66.1 66.0	111.8 111.7	25 24	140 139	$9.2 \\ 6.4$	7.3 5.0	346±11 344±10
5.	4/191	-	-	-	-	-	-	27	139	3.3	3.3	347±8

1 - "Named after 23rd Congress", 2 - "Sputnik", 3 - "Sytykan", 4 - "Yubileiny", 5 - summary determination for 4 pipes. Number of vectors involved in statistics: N - lumps of ore, n - samples. J, D - inclination and declination of the average direction of primary magnetization; k - clustering;  $\alpha_{95}$  - radius of confidence circle for 95% probability;  $\varphi,\lambda$  - coordinates of sampling site;  $\Phi,\Lambda$  - geographic latitude and longitude of the average paleomagnetic pole; A<sub>1</sub>, A<sub>2</sub> - semiaxes of ellipses of confidence for 95% probability, T.- paleomagnetic date.

The data were obtained on paleomagnetic poles of every pipe and general determination. According to paleomagnetic data the age of kimberlites is estimated as the Upper Devonian-Early Carboniferous, most probably Early Turnean (347+8 Ma), which agrees well with determinations by F.F.Brachfogel (1984, 1992) from a complex of radiological and other kimberlite dating evidence and is their independent confirmation and specification. Regarding the Khramov's magnetostratigraphic scale (1982), the dates obtained correspond to the middle part of the Tikhvinsk superzone of the Donetsk hyperzone characterized by a steble inverse polarity of geomagnetic field. The scatter of estimates for every pipe and the summary pole fall in a very narrow time span (3 Ma).

Zhitkov A.N. Report on topic 142391397 "Paleomagnetic determinations of age of Daldyn-Alakit kimberlites. VOSTSIBNIGGIMS, Irkutsk, 1994, 94 p. Zhitkov A.N. Study of paleomagnetism of kimberlite fields and

Zhitkov A.N. Study of paleomagnetism of kimberlite fields and hosting rocks of the Malo-Botuoba diamond-bearing region aimed at specifying the kimberlite pipe model. (Report, topic 407 for 1991-1992). VOSTSIBNIIGGIMS, Irkutsk, 1993, 77 p.

Savrasov D.I. Magnetism of Yakutian kimberlites. Thesis. Irkutsk, Earth's crust Inst., Irkutsk, 1978, 300 p.