### POLYBARIC MAGMA MIXING IN ALKALIC BASALTS AND KIMBERLITES: EVIDENCE FROM CORUNDUM, ZIRCON AND ILMENITE MEGACRYSTS

## Anthony J. Irving

Dept. of Geological Sciences, University of Washington, Seattle, WA 98195, U.S.A.

Megacryst phases enclosed in basanite, minette and kimberlite magmas provide a record of polybaric magmatic evolution during ascent of mantle-derived magmas. Evidence from trace element, isotopic and phase equilibrium data (e.g. Irving and Frey, 1984) implies that some megacrysts (e.g. Al-augite, pyropic garnet in basanites) could be cognate with the host magma, whereas others (e.g. corundum, zircon, anorthoclase in basanites) must be exotic to their hosts. One mechanism for the origin of the latter suite is precipitation from evolved magmas (e.g. nepheline benmoreites and phonolites) followed by mixing with later pulses of more primitive mafic magma. The possibility that some corundum megacrysts are xenocrysts from crustal metamorphic rocks, or are a product of desilication reactions between basanite, minette or kimberlite magmas and crustal rocks, must also be considered.

# Corundum and Zircon, Rubyvale, Queensland

A variety of megacryst species including corundum, zircon, anorthoclase, magnetite, ilmenite and pleonaste spinel occur within numerous Tertiary basaltic plugs in the Rubyvale region of central Queensland, Australia (Veevers et al., 1964; Stephenson, 1976), and adjacent alluvial deposits are actively mined for gem sapphire. The very similar geologic settings in southeast Asia (Barr and Macdonald, 1981) and Nigeria (Wright, 1970) imply that much of the world's gem sapphire forms megacrysts in alkalic basalts. The host rocks of four of the Rubyvale plugs (Mts. Leura, Pleasant, Hoy and Ball) are slightly to moderately evolved basanites (Mg/Mg+Fe<sup>2+</sup> 68.8-58.3). The corundum megacrysts (commonly 1-2 cm across) show multiple compositional zoning, especially in TiO<sub>2</sub> content (as reflected by zones of rutile exsolution). ICP bulk analysis of one megacryst shows minor contents of Fe, Ti (250 ppm), Mn (13 ppm), Cr (4 ppm), V (6 ppm), Sr (0.6 ppm) and Be (0.2 ppm).

Fluid inclusions within the corundums are very distinctive. They form planar arrays transverse to growth zoning, and heating-freezing studies reveal two main types. Type 1 is pure CO<sub>2</sub> (liquid + vapor). Such inclusions also occur in sapphires from minette at Yogo, Montana (Roedder, 1972). More common and associated with the first type are multi-phase inclusions composed of subequal amounts of CO<sub>2</sub> (liquid + vapor) and H<sub>2</sub>O containing halite, sylvite and other unidentified daughter minerals. Dissolution studies of the halide salts imply a high total salinity in the H<sub>2</sub>O phase ( $\sim$  35% NaCl and a similar value for KCl). These Type 2 inclusions remain unhomogenized at temperatures up to 685°C. Taken together the growth zoning, CO<sub>2</sub>-rich inclusions, and high inclusion homogenization temperatures suggest a high temperature (and pressure) magmatic origin for the corundum megacrysts. The occurrences of corundum intergrown with anorthoclase at Mt. Leura (Stephenson, 1976) and with sanidine + rutile in alkalic basalt from Loch Roag, Scotland (Upton et al., 1983) support this conclusion.

Zircon megacrysts from Rubyvale lack fluid inclusions. Two crystals analyzed by INAA contain relatively low bulk Hf (0.55-0.60%), U (7-30 ppm), Ta (0.7 ppm) and REE (see Fig. 1). These minor element abundances are quite similar to those in zircons from nepheline syenite pegmatites but much lower than those in zircons from granitoids (Kapustin, 1986; Murali et al., 1983). Taken with the occurrence of corundum in miascitic nepheline syenite pegmatites and the experimental evidence for instability of anorthoclase in basalt (Chapman, 1976), these data permit the inference that the Rubyvale megacryst assemblage may have crystallized from phonolitic magma at elevated (probably deep crustal) pressures. High pressure phonolites are known in north Queensland (Irving and Price, 1981) and have been modeled as products of fractional crystallization of basanites. The mechanism by which basanite magma would evolve to a condition of alumina-saturation is not clear, but may involve crystallization under moderate CO<sub>2</sub> and/or halogen partial pressures.

## Ilmenite Megacrysts in Kimberlites

Magnesian ilmenite megacrysts are well known in kimberlites. Preliminary results from INAA of an ilmenite from Monastery Mine, Orange Free State show relatively high contents of REE (0.3-0.5 x chondrites), Ta, Hf (28 ppm) and Sc (22 ppm). Furthermore the LREE abundances in this ilmenite are enriched by an order of magnitude over those expected for equilibrium with the host kimberlite (see Fig. 2), based upon known experimental partition coefficient values (McKay and Weill, 1976; Irving et al., 1978). In contrast data for an ilmenite megacryst from Benfontein are compatible with equilibrium with its host kimberlite (Fujimaki et al., 1984). The Monastery sample apparently crystallized from a liquid (or fluid) more enriched in LREE than the present host kimberlite thus implying magma mixing at this kimberlite pipe. The possibility that zircon megacrysts in kimberlites may have crystallized from evolved liquids and are exotic to their hosts will also require testing.

#### REFERENCES

BARR S.M. and MACDONALD A.S. 1981. Geochemistry and geochronology of the late Cenozoic basalts of southeast Asia. Bulletin of the Geological Society of America 92, pt. I, 508-512 and Pt. II, 1069-1142.

CHAPMAN, N.A. 1976. Inclusions and megacrysts from undersaturated tuffs and basanites, east Fife, Scotland. Journal of Petrology 17, 472-498.

- FUJIMAKÍ H., TATSUMOTO M. and AOKI K. 1984. Partition coefficients of Hf, Zr and REE between phenocrysts and groundmasses. Proceedings of the Fourteenth Lunar and Planetary Science Conference, Part 2, Journal of Geophysical Research 89, B662-B672.
- IRVING A.J. and FREY F.A. 1984. Trace element abundances in megacrysts and their host basalts: Constraints on partition coefficients and megacryst genesis. Geochimica et Cosmochimica Acta 48, 1201-1221.
- IRVING A.J. and PRICE R.C. 1981. Geochemistry and evolution of lherzolite-bearing phonolitic lavas from Nigeria, Australia, East Germany and New Zealand. Geochimica et Cosmochimica Acta 45, 1309-1320.
- IRVING A.J., MERRILL R.B. and SINGLETON D.E. 1978. Experimental partitioning of rare earth elements and scandium among armalcolite, ilmenite, olivine and mare basalt liquid. Proceedings of the Ninth Lunar and Planetary Science Conference, pp. 601-612. Pergamon.
- KAPUSTIN Y.L. 1986. Trace-element distribution in generations of accessory zircon in pegmatites. Doklady Earth Science Sections 277, 153-167.

McKAY G.A. AND WEILL D.F. 1976. Petrogenesis of KREEP. Proceedings of the Seventh Lunar Science Conference, pp. 2427-2447. Pergamon.

MITCHELL R.H. and BRUNFELT A.O. 1975. Rare earth element geochemistry of kimberlites. In Ahrens L.H. et al. eds., Physics and Chemistry of the Earth, Volume 9, pp. 671-685. Pergamon.

MURALI A.V. et al. 1983. Trace element characteristics, REE patterns and partition coefficients of zircons from different geological environments - a case study on Indian zircons. Geochimica et Cosmochimica Acta 47, 2047-2052.

- ROEDDER E. 1972. Composition of fluid inclusions. United States Geological Survey Professional Paper 440JJ, 164p.
- STEPHENSON P.J. 1976. Sapphire and zircon in some basaltic rocks from Queensland, Australia. Abstracts, 25th International Geological Congress 2, 602-603.
- UPTON B.G.J., ASPEN P. and CHAPMAN N.A. 1983. The upper mantle and deep crust beneath the British Isles: Evidence from inclusions in volcanic rocks. Journal of the Geological Society of London 140, 105-121.

VEEVERS J.J. et al. 1964. The geology of the Emerald 1:250,000 Sheet area, Queensland. Bureau of Mineral Resources, Geology and Geophysics, Report 68.

WRIGHT J.B. 1970. High pressure phases in Nigerian Cainozoic lavas: Distribution and geotectonic setting. Bulletin Volcanologique 34, 833-847.

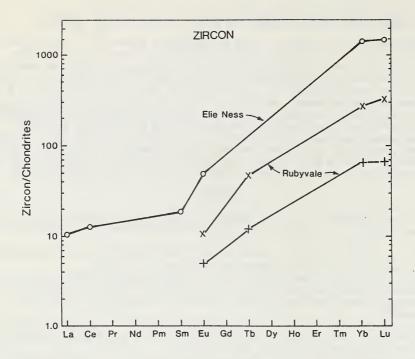


Figure 1. Chondrite-normalized REE data for zircon megacrysts from Rubyvale and Elie Ness, Scotland (Irving and Frey, 1984)

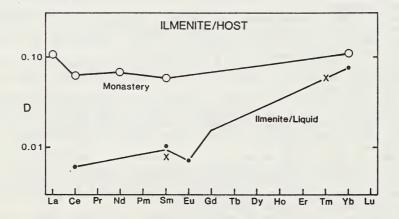


Figure 2. Comparison of ilmenite/host kimberlite REE abundance ratios (using host data from Mitchell and Brunfelt, 1975) with experimentally-determined D values.