

Geochronology and mantle source characteristics of kimberlites and related rocks from the Rae Craton, Melville Peninsula, Nunavut, Canada

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Introduction

Kimberlites of the Qilalugaq cluster were first discovered in 2003 by BHP Billiton approximately 10 km NE of the hamlet of Repulse Bay on the Rae Isthmus, southern Melville Peninsula, Nunavut, NE Canada. The Aviat kimberlite cluster was discovered by Stornoway Diamonds in 2002, approximately 60 km WSW of Igloolik on the Melville Peninsula (Fig 1). Subsequent detailed exploration program on these kimberlites reavealed that kimberlites in both of these clusters are diamondiferous. The grade of Aviat kimberlites (80-160 carat per hundred ton) is significantly higher than those of the Qilalugaq cluster (14-40 carat per hundred ton). In this study, we present the first U-Pb perovskite (CaTiO₃) geochronology of several kimberlite bodies to determine their emplacement ages along with the tracer isotope data to evaluate the evolution and nature of the mantle sources. We have also carried out new mineralogical and petrograpic characterisation of these kimberlites to comment on their nomenclature and degree of alteration.

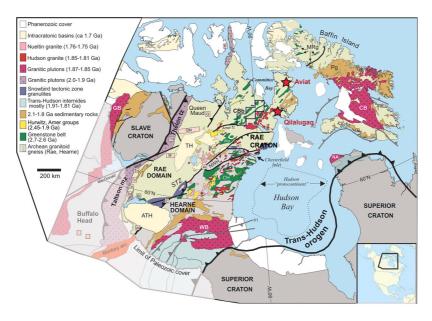


Fig 1: Simplified geological map of the Canadian Shield (modified after Berman, et al. (2005)). Location of Aviat and Qilalugaq kimberlite clusters shown by red stars.

Kimberlites and sampling

Both the Aviat and Qilalugaq kimberlite clusters intruded Archean tonalite-granodiorite gneisses, metasedimentary and metavolcanic rocks of the Rae craton (Dredge 2002). The Aviat cluster comprises 3 kimberlite pipes (AV1, AV4 and AV9) and five sheet-like bodies (AV2, AV3, AV5, AV8 and AV267; (Armstrong, et al. 2008)). All Aviat kimberlites occur within a 10 km long and 3

km wide corridor. The Aviat pipes mostly consist of coherent (hypabyssal) kimberlite (HK) to tuffisitic kimberlite breccia (TKB) while the sheet-like structures are dominated by HK and hypabyssal kimberlite breccia (HKB). The Qilalugaq cluster, located ~300 km to the south of Aviat, also consists of numerous kimberlite bodies including the Q1-4 complex that comprises four coalescing phases (A28, A48, A61 and A88) and several isolated kimberlite pipes (A34, A42, A59, A76, A94, A95, A152). Several dyke-like structures (Naujaat 1 to Naujaat 8) of variable width (1.5 m to 4.5 m) were also discovered (Kupsch and Armstrong 2013). The dominant phase in most of the Qilalugaq pipes and the dykes is a massive volcanoclastic kimberlite (MVK) that can be classified as TKB, which commonly transitions toward HK with depth. They all show variable degrees of country rock dilution with the dykes being diluted more than the pipes. Qilalugaq kimberlites contain peridotitic xenoliths with mutli-modal Re-depletion ages ranging from Archean (3.0-2.7 Ga), Paleoproterozoic (2.1–1.7 Ga) and "Recent" (<1 Ga), suggesting the kimberlites intruded through a complex layered lithosphere (Liu, et al. 2016).

Samples were collected from both kimberlite clusters in the summer of 2012, from the pits and trenches remaining after the earlier exploration. Some drill core samples were also used for petrographic and mineralogical study. Relatively hard and unaltered HK were chosen for perovskite recovery following a standard heavy mineral separation.

Petrography and Mineralogy

The Aviat pipes and sheets examined are all hypabyssal (coherent) rocks and are broadly similar from a mineralogical perspective. Large olivine grains are set in a groundmass of olivine, apatite, perovskite, spinel (chromite zoned to Ti-magnetite), ilmenite, calcite and poikilitic, zoned phlogopite. All olivine grains are variably replaced by calcite and/or serpentine. However, the Aviat sheets are distinctive in that apatite, perovskite and calcite are less abundant, and there is a higher proportion of groundmass ilmenite and spinel as compared to samples from the Aviat pipes. In the sheets, rutile +/dolomite is also observed as secondary mineral(s). The Quilalugaq and Naujaat samples are quite variably altered; the Naujaat dykes being extensively carbonatized (>60% dolomite, +/-calcite), or silicified (45-60 wt% SiO₂). In the Naujaat samples, the olivine is completely replaced; only groundmass apatite, spinel and phlogopite are preserved, with secondary rutile and dolomite. Hypabyssal (coherent) samples from A34, A88 and A94 are similar, with olivine (replaced by serpentine and/or carbonate) set in a groundmass of apatite, spinel (chromite zoned to Ti-magnetite), phlogopite, calcite (+/-dolomite), +/-perovskite. All of the samples noted above from Aviat, Quilalugaq and Naujaat classify as being kimberlite, using the criteria of Woolley, et al. (1996). However, hypabyssal samples from Quilalugaq A61 consist of olivine set in a groundmass of apatite, perovskite, spinel, tetraferriphlogopite, calcite, augite and Ti-andradite (schorlomite-melanite) garnet; the latter two minerals consistent with these samples being ultramafic lamprophyres or orangeites, but not Gp I (archetypal) kimberlites. Mineral chemistry studies are currently being undertaken to refine the classification of these rocks.

Results and Discussion

ID-TIMS U-Pb analyses of seven perovskite fractions from both Aviat and Qilalugaq kimberlites indicate their early Cambrian age. While the major pipes from both clusters were emplaced between 531 and 536 Ma, the sheet-like kimberlite bodies at Aviat are ~30 Ma younger, emplaced between 503 and 507 Ma. Perovskite U-Pb results from the Naujaat dykes, gave a Proterozoic date, but is considered unreliable due to their unusually high common Pb correction. In general, Cambrian kimberlites are uncommon worldwide. The ~530 Ma Venetia kimberlite in South Africa is perhaps the best known. However, there are several kimberlite clusters of Late Neoproterozoic to Eocambrian age that have been reported from North America and some of these have economic diamond potential. These include Pelly Bay from Rae craton (Kienlen, et al. 2008; Liu, et al. 2016), Snap Lake and Gahcho Kué from SE Slave craton (Heaman, et al. 2004), Renard and Wemindji from the Superior craton (Tappe, et al. 2016), Chicken Park and George Creek from the Wyoming craton (Heaman, et al. 2004). Several other kimberlites (Labrador Sea Province, Lac Beaver kimberlite) and related rocks

(ultramafic lamprophyres and carbonatites from Labrador and west Greenland) of similar age have been reported from eastern North America and western Greenland (Heaman, et al. 2004; Tappe, et al. 2014). This region of north-eastern North America and western Greenland has been termed the Eocambrian/Cambrian Labrador Sea Province and the magmatism here has been suggested to be associated to Eocambrian rifting along the northeastern margin of Laurentia and opening of the Iapetus Ocean (Heaman, et al. 2004).

Sr-Nd tracer isotopes were also obtained on the same perovskite fractions used for U-Pb dating from Qilalugaq and Aviat. The main kimberlite pipes at Qilalugaq (A61 and A94) and Aviat (AV1) have overlapping, relatively restricted ⁸⁷Sr/⁸⁶Sr_i, varying between 0.70405 and 0.70465. The Sr isotopic signature of these pipes are slightly more radiogenic than typical Cretaceous age Group I kimberlites from South Africa (Nowell, et al. 2004) but are very similar to other Canadian kimberlites, such as Lac de Gras, Kirkland Lake, Somerset Island, and Fort à la Corne (e.g., Tappe, et al. (2013)). In contrast, the sheet-like bodies (AV2 and AV6) have considerably more radiogenic signatures (87Sr/86Sr_i = 0.70948-0.70969). Initial εNd values from Qilalugaq kimberlites define a very narrow range of +4.2 and +4.7, consistent with the restricted ⁸⁷Sr/⁸⁶Sr_i values. In contrast, those from Aviat kimberlites cluster around chondritic values (-1.5 to +1.5). The homogeneous Sr-Nd isotopic nature of the main kimberlites from Aviat and Qilalugaq and their position on a mantle array in εNd vs ⁸⁷Sr/⁸⁶Sr space suggest their derivation from the asthenosphere or transition zone. However, the more radiogenic signature recorded by the perovskites from sheet-like intrusions at Aviat are unlikely to be due to any contamination/alteration, as perovskite is much more robust to the effects of crustal contamination. Hence, these younger kimberlites may be derived from a Subcontinental Lithospheric Mantle source, similar to that recorded in high-T peridotite xenoliths entrained in the nearby Somerset Island kimberlites (Schmidberger and Francis 2001).

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