

## Diamonds from the Guaniamo area, Venezuela

Kaminsky, F.V.<sup>1</sup>, Zakharchenko, O.D.<sup>2</sup>, Channer, D. M. DeR.<sup>3</sup>, Blinova, G.K<sup>2</sup>, and Maltsev, K.A.<sup>4</sup>

1. KM Diamond Exploration, 815 Evelyn Drive, West Vancouver, B.C. V7T 1J1, Canada, Tel. & Fax (1-604) 925-8755, E-mail: kaminsky@fox.nstn.ca

2. Institute of Diamonds, Russian Academy of Natural Sciences, 155-5/10 Litovskii Blvd., Moscow 117593, Russia

3. Guaniamo Mining Company, Centro Gerencial Mohedano, Cruce Av. Mohedano, Urb. La Castellana, Caracas, Venezuela

4. Vernadsky Institute of Geochemistry and Analytical Chemistry, Russian Academy of Sciences, Moscow, Russia

Diamondiferous placer deposits were discovered in the Guaniamo area in 1968, and since then approximately 20-25 million carats have been produced from this area. Stones up to 40 and 60 carats in weight have been reported here. Recently diamondiferous kimberlite sills were discovered in the Quebrada Grande, Guaniamo River right tributary valley which are the primary sources of the diamonds. We have studied over 4,000 diamonds from both, kimberlites and placer deposits.

**Morphology.** Diamonds from the Guaniamo area include dodecahedroids, octahedra, octahedron + dodecahedroid (O-D) combination-type crystals and their twins and aggregates. In all sills and placer deposits, dodecahedroids are predominant, accounting for about 45-50% of the respective samples. Combination-type O-D crystals are also present in rather high proportions, sometimes up to 23% of the samples. All *dodecahedral* diamond crystals are rounded dodecahedroids. Eight groups of them were distinguished based on the nature of the dominant crystal surface. Dodecahedroids with concentric striation are predominant. They compose 60-70% of all dodecahedral diamonds. *Octahedral* diamonds are subdivided into the following three groups: (1) Octahedra with stepwise lamellar development of trigonal faces, (2) Octahedra with clearly concentric development of crystal faces and (3) Octahedra with bitrigonal faces. Diamonds of the last type are commonly predominant. Single *cubic-habit* crystals were found. They are represented by tetrahexaedroids and combination-type crystals. There are many varieties of *twinned and aggregated diamonds*. Their proportion in the larger diamonds increases. Macles, or octahedral spinel-type twins are most common. Only a small number of diamonds from the Quebrada Grande placer show post-magmatic mechanical erosion features and, may therefore have been transported over some distance.

By **color**, colorless diamonds are predominant: they comprise 40-50% of all diamonds, but in some placers form up to 70% of stones. Smoky-brown, grey and green stones account for 10-20% each of all stones. 60-80% of diamonds have *pigmentation spots*. These are typically isolated spots that do not affect the overall coloration of the crystal. Rarely, these spots occur in clusters. In the overwhelming majority of cases, the spots are bright green. However, several diamonds showing brown pigmentation spots were found in the Quebrada Grande placer collection.

All the sills studied have a marked predominance of diamonds showing blue **photoluminescence** (PL) and a rather high proportion of crystals exhibiting no visible luminescence. In general, PL characteristics of diamonds from the Guaniamo sills are identical to those of diamonds from the Guaniamo placer deposits. 55.8% and 57.6%, respectively, of the diamonds in these two groups,

have blue luminescence. The diamonds from the sills and placers also show similar abundances of stones with yellow-green and pale pink luminescence, of diamonds with uncertain PL color and, of diamonds with no visible luminescence. Diamonds from these two sources which exhibit inhomogeneous luminescence show only very small differences. Hence, diamonds from the sills and the placers are essentially identical in their luminescence characteristics, and the sills are therefore interpreted as representing the major source of Guaniamo placer diamonds.

**Diamond IR spectra** from the sills and placers are rather similar. They all show bands representing IaA, IaB and P (platelets) nitrogen impurities and do not show resolvable concentration of single nitrogen atom impurities of type Ib. Type II diamonds were not found in the samples studied. All Guaniamo diamonds studied belong to the transitional IaAB type. More than half the grains show significant amounts of hydrogen structural impurities, while the remainder appear to contain moderate amounts. The majority of diamonds show a predominance of IaB-type nitrogen impurities compared to IaA impurities, though in a few crystals, IaA nitrogen impurities are more abundant than IaB impurities. As much as 20% of all diamonds studied contain relatively low ( $5 \times 10^{19}$  atoms/cm<sup>3</sup>) concentrations of all three types (IaA, IaB and B) of nitrogen impurities. Concentrations of IaA nitrogen impurities and hydrogen impurities are not very high. Their range in concentration are  $0.6\text{--}17.3 \times 10^{19}$  atoms/cm<sup>3</sup> and  $0.3\text{--}3.5 \text{ cm}^{-1}$ , respectively. IaB nitrogen impurities are typically more abundant than IaA impurities, with concentrations ranging from 1 to  $4 \times 10^{19}$  atoms/cm<sup>3</sup>. Concentrations of B nitrogen impurities range from 0.4 to  $10.8 \times 10^{19}$  atoms/cm<sup>3</sup>, commonly approaching IaA impurity concentrations. The predominance of IaB and P nitrogen impurities over IaA impurities in the diamonds studied suggests that Guaniamo area diamonds are characterized by higher temperatures of formation than diamonds from Siberian pipes.

The total range of variation in **carbon isotopic compositions** of Guaniamo diamonds analyzed is from  $\delta^{13}\text{C} = -3.2\text{‰}$  to  $-28.7\text{‰}$ . The majority of diamonds are isotopically light, i.e., show  $\delta^{13}\text{C} < -10\text{‰}$  (from  $-10.1$  to  $-28.7\text{‰}$  for 95% of diamonds).

Guaniamo district diamonds contain both eclogitic- and ultramafic-suite **mineral inclusions**. Diamonds containing eclogitic-suite inclusions are predominant. For example, only 10% of garnets belong to the ultramafic association. The finer the diamond size, the greater is the proportion of eclogite-type inclusions.

Eclogitic (E-type) garnet of the *pyrope-almandine* series have rather high CaO content in most of the garnets (up to 10-13%) and do not have high Na<sub>2</sub>O content (0-0.26%). *Omphacite* clinopyroxene, like orange garnet, occurs relatively commonly as inclusions in the diamonds studied. In some cases, composite pyroxene and garnet inclusions are observed. Na<sub>2</sub>O content varies from 3.63% to 7.46 wt. %, which correspond to 25 to 51 mol. % jadeite component. Al<sub>2</sub>O<sub>3</sub> content varies from 7.21 to 13.53 wt. %, and positively correlated with Na<sub>2</sub>O. In addition, pyroxenes included in diamonds from the Guaniamo area exhibit rather high concentrations of Na<sub>2</sub>O. Diamonds that contain orange garnet and omphacite inclusions also commonly enclose colorless mineral inclusions which considering the E-type paragenesis of these diamonds, are probably formed by *coesite*. The *rutile* inclusions have compositions which are typical of E-type rutile inclusions in diamonds. *Ilmenite* inclusions were found in two diamonds. Ilmenite in diamonds from other diamondiferous areas worldwide are commonly very rich in both MgO and Cr<sub>2</sub>O<sub>3</sub>. A notable feature of the ilmenite inclusions from the Guaniamo area studied here is their rather low MgO content, from 0.91 to 1.04 wt. %. Mg-poor ilmenite inclusions were previously

found in only two other diamonds, from Brazil by Meyer and Svisero [1975]. The ilmenites studied are also notably characterized by high MnO content (about 1%). *Sulphides* were also found included in some Guaniamo area diamonds. Based on their compositions, in particular their low Ni content, they are interpreted to be eclogitic.

Ultramafic (UM-type) inclusions include Cr-garnet, chrome spinel and olivine. *Olivine* inclusions in diamonds from the Guaniamo area have 93-94.5% forsterite contents and contain insignificant amounts of Cr<sub>2</sub>O<sub>3</sub> (0.04-0.08 wt. %), CaO (0.01 wt. %) and NiO (0.33-0.39 wt. %). The *chromite* inclusions have Cr<sub>2</sub>O<sub>3</sub> and Al<sub>2</sub>O<sub>3</sub> contents ranging from 63.6 to 66 wt. % and 6.76 to 7.25 wt. %, respectively, which is typical for chromite inclusions in diamonds from the majority of diamond deposits worldwide. All of the *pyrope* inclusions are high-Cr knorringite-bearing garnets (6.9 to 13.8 wt. % Cr<sub>2</sub>O<sub>3</sub>) with low iron contents (13.5-15.0 wt. % FeO). One of the Cr-pyropes studied appears to be Ca-poor, and the other three have rather high Ca-contents. Two of them fall within the lherzolite field, which is somewhat atypical for pyrope inclusions in diamonds. The other two garnets fall within the diamond-association field, with one having rather low Ca for this field. One diamond contains a *titano-magnetite* inclusion. In another diamond, a single notable occurrence of included *magnesio-wustite* was observed. It has the following composition: 0.15 wt. % Cr<sub>2</sub>O<sub>3</sub>; 19.42 wt. % FeO; 80.12 wt. % MgO.

Four diamonds have composite inclusions of coexisting garnet and clinopyroxene. Equilibrium temperatures and pressures for these mineral pairs lies within the range 1186-1218 °C temperature and 41.4-61.5 kb pressure. These results are consistent with results of the IR studies that Guaniamo diamonds have higher pressure-temperature conditions of formation than diamonds from most of the other diamondiferous areas worldwide.

Diamonds from different sills are distinct in most of their properties. This distinctiveness of diamonds from each of the sills can be used as a "diamond fingerprinting" tool in prospecting.

## References

Baptista, G.J., and Svisero, D.P., 1978, Geologia de los depositos diamantiferos de la parte norooccidental de la Guayana Venezolana: Republica de Venezuela, Ministerio de Energia y Minas, Direccion General Sectorial de Minas y Geologia, XIII, No. 24 (Agosto, 1978), p. 3-46.

Meyer, H.O.A., and Svisero, D.P., 1975, Mineral Inclusions in Brazilian Diamonds: Physics and Chemistry of the Earth, 9, p. 785-795.

Nixon, P.H., 1988, Diamond source rocks from Venezuela: Indiaqua, 51, p. 23-29.

Nixon, P.H., Griffin, W.L., Davies, G.R., and Condcliffe E., 1995, Cr garnet indicators in Venezuela kimberlites and their bearing on the evolution of the Guayana craton: Kimberlites and related rocks, 1, Brasilia, p. 378-380.

Svisero, D.P., and Baptista, G.J., 1973, Inclusiones en los diamantes de la Quebrada Grande, Distrito Cadeno, Estado Bolivar, Venezuela: II Congr. Lat. Amer. Geol., Caracas, 1973, Resumenes, p. 158-160.